

1 **TITLE:** Spatial-temporal interpolation of satellite geomagnetic data to study long-distance
2 animal migration

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34 **Abstract**

35 Introduction: Increased access to remote sensing datasets presents opportunities to model an
36 animal's in-situ experience of the landscape to study behavior and test hypotheses such as
37 geomagnetic map navigation. MagGeo is an open-source tool that combines high spatiotemporal
38 resolution geomagnetic data with animal tracking data. Unlike gridded remote sensing data,
39 satellite geomagnetic data are point-based measurements of the magnetic field at the location of
40 each satellite. MagGeo converts these measurements into geomagnetic values at an animal's
41 location and time. The objective of this paper is to evaluate different interpolation methods and
42 data frameworks within the MagGeo software and quantify how accurately MagGeo can model
43 geomagnetic values and patterns as experienced by animals.

44 Method: We tested MagGeo outputs against data from 109 terrestrial geomagnetic observatories
45 across 7 years. Unlike satellite data, ground-based data are more likely to represent how animals
46 near the Earth's surface experience geomagnetic field dynamics. Within the MagGeo framework,
47 we compared an inverse-distance weighting interpolation with three different nearest-neighbor
48 interpolation methods. We also compared model geomagnetic data with combined model and
49 satellite data in their ability to capture geomagnetic fluctuations. Finally, we fit a linear mixed-
50 effect model to understand how error is influenced by factors like geomagnetic activity and
51 distance in space and time between satellite and point of interest.

52 Results and conclusions: The overall absolute difference between MagGeo outputs and
53 observatory values was less than 1% of the total possible range of values for geomagnetic
54 components. Satellite data measurements closest in time to the point of interest consistently had
55 lowest error which likely reflects the ability of the nearest neighbour in time interpolation method
56 to capture small continuous daily fluctuations and larger discrete events like geomagnetic storms.

57 Combined model and satellite data also capture geomagnetic fluctuations better than model data
58 alone across most geomagnetic activity levels. Our linear mixed-effect models suggest that most
59 of the variation in error can be explained by location-specific effects originating largely from local
60 crustal biases, and that high geomagnetic activity usually predicts higher error though ultimately
61 remaining within the 1% error range. Our results indicate that MagGeo can help researchers
62 explore how animals may use the geomagnetic field to navigate long distances by providing access
63 to data and methods that accurately model how animals moving near the Earth's surface experience
64 the geomagnetic field.

65

66 **Keywords:**

67 Geomagnetism, navigation cues, animal movement, wildlife tracking, Swarm constellation,

68 CHAOS-7 model

69 **1. Introduction and background**

70 *1.1 Introduction*

71 Understanding how migratory animals navigate the landscape is challenging not least
72 because of the spatiotemporal range of some migrations (Wilcove & Wikelski, 2008). Access to
73 remote sensing imagery has influenced our understanding of how and why an animal interacts with
74 its environment (Pettorelli et al., 2014). However, the predominant use of optical remote sensing
75 imagery often restrains how we model an animal's relationship to its surroundings. Remote sensing
76 satellites with sensors, such as Synthetic Aperture Radars (SARs) and Light Detection and
77 Ranging (LiDAR), offer opportunities for novel lines of questioning in wildlife movement
78 ecology. Satellites with geophysical sensors measuring the Earth's magnetic field are an example
79 of an underexplored non-optical resource that can bring new insights, especially with regards to
80 the magnetic map hypothesis (Lohmann et al., 2007; Mouritsen, 2014; Naisbett-Jones et al., 2017;
81 Wiltschko & Wiltschko, 2013).

82 The geomagnetic field is an invisible shield around the Earth that protects the planet from
83 incoming solar particles (Campbell, 2003). Large scale geomagnetic patterns vary predictably
84 across space and time, thus allowing human navigators to reliably use geomagnetic information
85 for wayfaring for many centuries. Animals who are capable of sensing and perceiving the
86 geomagnetic field may also use geomagnetic patterns to make navigation decisions during
87 migration (Lohmann et al., 2007; Mouritsen & Heyers, 2016; Wiltschko & Wiltschko, 2021). The
88 underlying mechanisms of the geomagnetic navigation strategies vary between species and are
89 highly debated in the literature (Åkesson et al., 2005; Holland et al., 2009; Muheim et al., 2018;
90 Pollonara et al., 2015). Specifically, migratory navigation consists of two tasks: 1) knowing the
91 current location (geographic positioning) and 2) knowing in which direction to go (compass

92 orientation). Some research suggests that animals use geomagnetic information for orientation,
93 and it is also possible that animals use two or more geomagnetic values to build cognitive maps
94 for positioning (although this has not been proven and is at present still debated) (Mouritsen, 2018;
95 Sokolovskis et al., 2018; Wikelski et al., 2015). Physiological capabilities to sense geomagnetic
96 values have been tested in laboratory experiments, alongside physical or virtual displacement that
97 demonstrate how a bird's migratory direction oscillates with changes in a magnetic environment
98 (Kishkinev, 2015). The exact sensitivity range to changes in absolute geomagnetic values is
99 unclear and likely varies by species, internal states, and external conditions. Some experiments
100 suggest ranges from 15 nT to 200 nT for total intensity (Beason & Semm, 1987; Semm & Beason,
101 1990), 2° to 5° for inclination (Schwarze et al., 2016), and at least 8° for declination sensitivity
102 (Chernetsov et al., 2017). There are even fewer experiments that have explored how wild migrants
103 respond to the geomagnetic field during the actual migration and what strategies they use for
104 orientation and positioning. To understand what happens outside of controlled laboratory settings,
105 there has been a push to test geomagnetic strategies from a data-driven, geospatial perspective by
106 taking advantage of open-source geomagnetic models and satellite data (Zein et al., 2021, 2022).

107 To look at this, previous studies have successfully combined geomagnetic model estimates
108 with animal tracking data (Åkesson & Bianco, 2016, 2017; Sokolovskis et al., 2018; Zein et al.,
109 2021). Model estimates are typically geomagnetic field values predicted using a set of model
110 coefficients that are informed by satellite geomagnetic data collected during periods of low
111 geomagnetic activity otherwise known as quiet-time (Chulliat et al., 2015; Matzka et al., 2010;
112 Olsen et al., 2006). Model estimates are continuous across time allowing daily estimates for any
113 latitude, longitude, and altitude combination (location in 3D space). As such, estimates are useful
114 for movement ecologists trying to understand how geomagnetic information can influence animal

115 behavior across the entire migratory trajectory. For example, Åkesson & Bianco (2016, 2017) used
116 the 11th Generation International Geomagnetic Reference Field (IGRF-11) model to create
117 simulated migratory paths built from model estimates. They compared these trajectories with
118 observed paths recorded by migratory birds carrying GPS trackers. Zein et al. (2021) used IGRF-
119 12 to combine migratory bird tracks with model estimates to test different geomagnetic navigation
120 strategies. Studies concluded that a geomagnetic compass mechanism is possible though further
121 exploration about an animal's response to actual geomagnetic conditions during migration was
122 restricted due to data limitations (Zein et al., 2022).

123 Model estimates capture much of the variability in the geomagnetic field, but not all and
124 especially not the dynamics that may affect an animals' instantaneous responses to the
125 contemporaneous geomagnetic conditions. The geomagnetic field varies across space and time at
126 different scales. Across space, there are both large planetary variations of the field generated by
127 Earth's core and small-scale changes related to crustal field generated by magnetic rocks in the
128 Earth's crust. Temporally, the crustal field changes slowly over millions of years, while the core
129 field changes over years to decades – this is called secular variation. However, the field also
130 changes over the course of the day in response to the variable solar wind, which generates
131 fluctuations in the ionosphere and magnetosphere (Courillot & Le Mouel, 1988). Solar storms and
132 large solar flares can further lead to disturbances over much shorter temporal scales (seconds to
133 hours) known as geomagnetic storms, whose effects can range from benign and beautiful auroral
134 light displays to technological disruptions, such as satellite anomalies and power blackouts
135 (Babayev et al., 2006; Hapgood, 2012; Kikuchi, 2003; Lanza & Meloni, 2006). Specifically,
136 model estimates omit fine-scale spatial variability created by very local but acute geomagnetic
137 anomalies in the crustal field, and they also do not represent the short-term temporal dynamics of

138 the field, as the models are derived from data largely measured during quiet-time conditions. This
139 means that model estimates do not wholly represent the geomagnetic landscape as experienced by
140 animals (Beggan, 2022; Benitez-Paez et al., 2021; Bianco et al., 2019; Krylov, 2017; Zein et al.,
141 2021).

142 Raw geomagnetic data collected continuously by sensors on-board satellites are a source
143 of localized higher spatial and temporal resolution geomagnetic information. The most
144 comprehensive example is the European Space Agency's (ESA) recent mission of Swarm satellites
145 (European Space Agency, 2020; Friis-Christensen et al., 2006; Olsen et al., 2013). Since 2014,
146 two satellites in near-polar parallel circular orbits and a third in a drifting local time circular orbit,
147 are continuously collecting data as they move over the Earth's surface at an altitude of 450-510
148 km (Friis-Christensen et al., 2006, 2008). Unlike model estimates, satellites measure the actual
149 magnetic conditions, which include contributions from all major magnetic sources (core, crust,
150 ionosphere) as well as the real-time effects of the interaction with the solar wind. Swarm data are
151 openly available through the VirES interface (European Space Agency, 2021; Kloss, 2021).

152 MagGeo is an open-source tool that takes advantage of the high resolution of Swarm data
153 and combines it with model estimates to create an accurate representation of magnetic conditions
154 at a specific location and moment in time (Benitez-Paez et al., 2021), thus enabling linkage of
155 satellite geomagnetic data with animal tracking data (for example, trajectories collected by GPS
156 tags). MagGeo gets model estimates from the 7th Generation of the CHAMP, Ørsted and SAC-C
157 (CHAOS-7) model of the Earth's magnetic field (Finlay et al., 2020). A major challenge when
158 combining animal tracking data with environmental variables like geomagnetic data is matching
159 the spatial and temporal resolution of different datasets (Brum-Bastos et al., 2021). Interpolation
160 methods are required to overcome these differences because it is rare that a measured

161 environmental variable and a moving animal coincide perfectly in space and time. After correcting
162 for the difference in altitudes between satellite orbits and animals moving near the Earth's surface,
163 MagGeo uses an inverse-distance weighting interpolation method to combine Swarm satellite data
164 with GPS tracking data to allow movement ecologists to test hypotheses about geomagnetism and
165 animal movement from a geospatial data-driven perspective.

166 Benitez-Paez et al. (2021) performed an initial error and accuracy analysis of MagGeo
167 though, it was limited to three test locations in Europe for six days of variable geomagnetic activity.
168 A more thorough error and accuracy analysis is required to ensure MagGeo's useability for
169 locations outside of Europe and across various time periods. Furthermore, MagGeo assumes that
170 (1) inverse-distance weighting and (2) a combination of model estimates and satellite data are the
171 most likely interpolation method and data structure to accurately model the geomagnetic field as
172 experienced by animals. These assumptions have not been tested.

173 We perform a global error and accuracy analysis for MagGeo by testing more than 100
174 locations across 7 years (2014-2020) and following common practices outlined by the geophysical
175 community (Beggan, 2022; Chulliat et al., 2015; Macmillan & Olsen, 2013). We evaluate accuracy
176 measures across four spatiotemporal interpolation methods: (1) inverse-distance weighting, and
177 three nearest-neighbor methods for (2) space, (3) time and (4) spacetime. Next, we compare the
178 accuracy between model estimates (from CHAOS-7) and fused model estimate and satellite data
179 (combining CHAOS-7 and Swarm satellite data). We highlight important considerations for
180 researchers hoping to model the geomagnetic field through MagGeo to ask questions about animal
181 navigation. We also demonstrate the benefits of performing error and accuracy assessments of
182 remotely sensed environmental data that are applicable to movement ecologists. We believe that
183 studies like ours encourage cross-disciplinary collaboration and will become increasingly

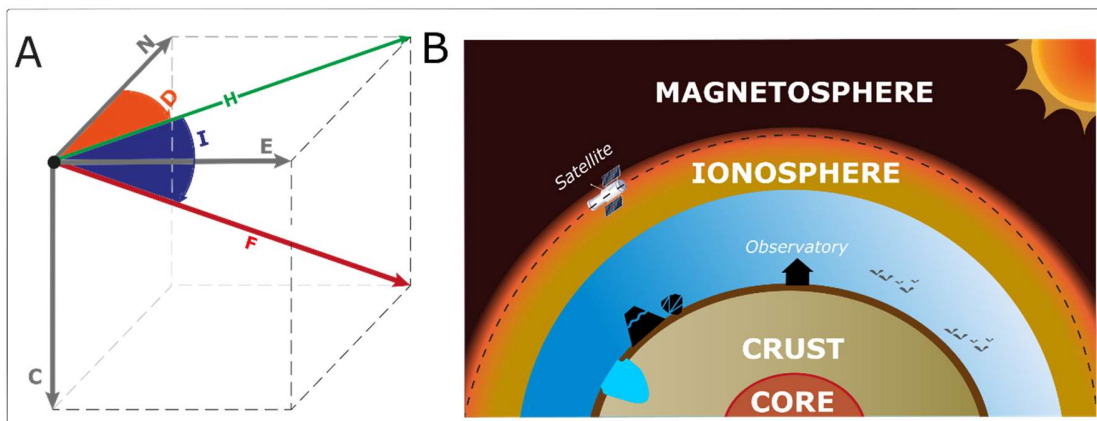
184 important with the current trends in technology evolution and data accessibility (Guilford et al.,
185 2011; Kays et al., 2022; Nathan et al., 2022).

186 *1.2 Background*

187 *1.2.1 Earth's geomagnetic field*

188 The Earth's magnetic field is notionally like a bar magnet with field lines exiting the
189 geomagnetic south pole (near Antarctica) and entering the geomagnetic north pole (presently
190 within Ellesmere Island). In detail, the geomagnetic field is far more complex and has various
191 components apart from polarity (Figure 1A). Total field intensity (F) and the horizontal component
192 of the field intensity (H) are two scalar quantities that measure the magnitude of the geomagnetic
193 field vector in nanoteslas (nT). Inclination (I) and declination (D) are angular components of the
194 geomagnetic field vector measured in degrees. Inclination refers to the angle between the field
195 vector and the Earth's horizon whereas declination is the angle between magnetic north and the
196 geographic north pole. Declination is used to align the geomagnetic field on the Earth and is not a
197 natural property of the field since it requires additional knowledge of the relative position of the
198 geographic North and South poles. Geometrically, these components (FHID) can be calculated
199 from values collected by geomagnetic sensors which are measured in the North (N), east (E) and
200 center (C) cartesian coordinate system (Figure 1A) (Campbell, 2003). There are multiple sources
201 of Earth's geomagnetic field, the principal being the geodynamo in the liquid outer core which
202 accounts for around 98% of the total field and has a surface strength of between 23000-60000 nT.
203 Next, the magnetic minerals in the local subsurface (crust) varies between 10-1000 nT depending
204 on location. Electrical currents in the ionosphere at approximately 100-1000 km above the Earth's
205 surface and electrical currents in the magnetosphere which extends even further into outer space
206 are the two final sources of the geomagnetic field (Figure 1B). Different altitudes at the same

207 geographic coordinate will have different geomagnetic values depending on the proximity to the
 208 geomagnetic sources (Campbell, 2003; Hulot et al., 2010; Thébault et al., 2010). The typical
 209 strength of the external field in magnetically quiet conditions is 20-50 nT but rises to >1000 nT in
 210 active periods. Geomagnetic field activity is quantified on a quasi-logarithmic scale called the Kp
 211 index (with values of 0-9) which often accompanies open-source geomagnetic data (Matzka et al.,
 212 2021). During periods of high solar activity such as geomagnetic storms, values of the geomagnetic
 213 field can change rapidly particularly at the mid to high latitudes.



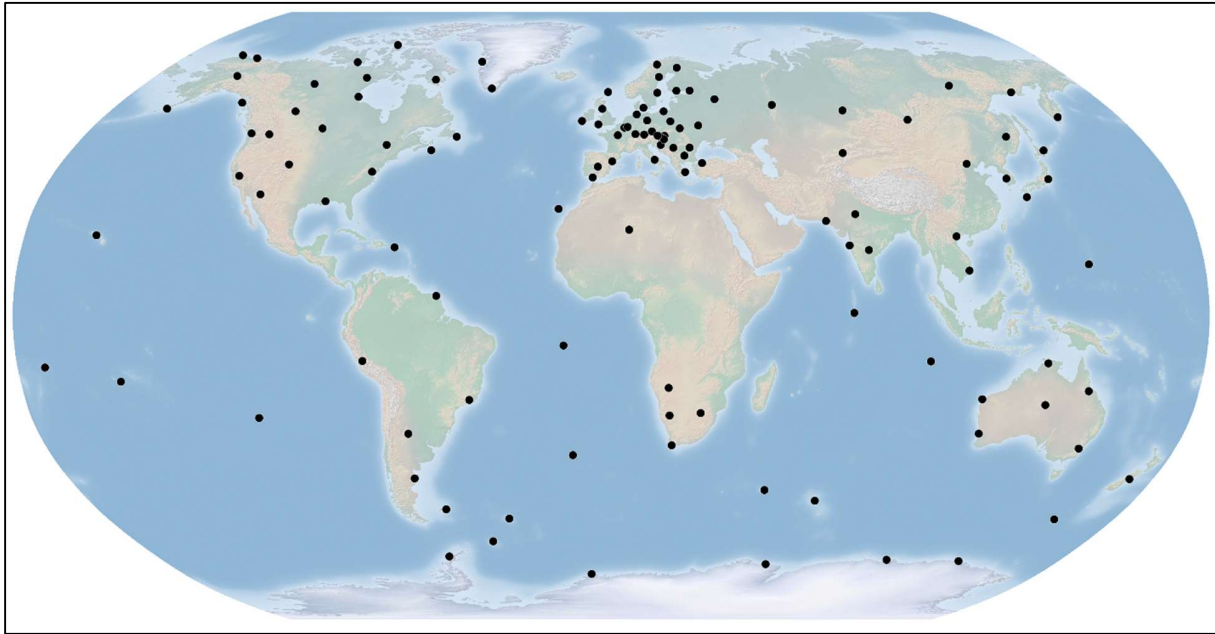
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215 **Figure 1.** The main components of the geomagnetic field. **A.** The geomagnetic coordinate system is in the
 216 North-East-Centre (NEC) coordinate frame shown in gray with the 4 geomagnetic components highlighted
 217 in color: declination (D) in orange, inclination (I) in blue, horizontal intensity (H) in green, and total
 218 intensity (F) in red. **B.** The 4 main contributors to the geomagnetic field from innermost to outermost: core,
 219 crust, ionosphere, and magnetosphere. Examples of geomagnetic anomalies due to lithosphere composition
 220 represented by symbols for water body, volcano, and exposed magnetic rock. Satellite orbiting in the
 221 ionosphere and the geomagnetic observatory are representations for the two main geomagnetic data sources.

222 1.2.2 Geomagnetic data sources

223 Geomagnetic data are traditionally collected at ground-based observatories. The
 224 INTERMAGNET network of observatories (Figure 2) currently has 126 operational stations across
 225 the world that collect geomagnetic data (INTERMAGNET, 2020) available at second-, minute-
 226 and hour cadences. Since observatories are located at ground level, their data are heavily
 227 influenced by the core and crustal components of the geomagnetic field (Thébault et al., 2010).

228 While they have high temporal resolution, data from INTERMAGNET observatories are limited
229 to their locations which are irregularly distributed. For example, there are only six stations in
230 Africa and six stations in South America. The low station density impairs study of long-distance
231 animal migration that can span multiple continents.



232
233 **Figure 2.** Global distribution of INTERMAGNET observatories (n=126).

234 In contrast to ground stations, polar-orbiting satellites with on-board magnetometers
235 collect globally distributed data on the geomagnetic field for locations on their orbit. These
236 satellites collect data at high altitudes (400-500 km) and are strongly influenced by the ionosphere
237 (Benitez-Paez et al., 2021; Campbell, 2003) (Figure 1B). An ongoing mission to gather satellite
238 geomagnetic data is operated by ESA with their launch of three Swarm satellites in late-2013
239 (Friis-Christensen et al., 2006; Olsen et al., 2013). One-second resolution data from these satellites
240 are available within 96 hours of collection and can be accessed through the VirES platform
241 (European Space Agency, 2021).

242 A combination of observatory data, satellite data, and ground data are used to inform
243 creation of geomagnetic models which are often spherical harmonic models determined by a set
244 of so-called Gauss coefficients (Chulliat et al., 2015; Olsen et al., 2006; Sabaka et al., 2020).
245 Geomagnetic values are then estimated from these model coefficients and are used for geophysical
246 studies, long-term monitoring, resource exploration and extraction, and are updated periodically
247 to account for the non-linear continuous changes in the geomagnetic field (secular variation). Due
248 to the ease at which model estimates can be computed for each unique location in 3D space, they
249 are often used in non-geophysical field applications, and have previously been used in the analysis
250 of animal migrations (Boström et al., 2012; Komolkin et al., 2017). Due to the complexity of the
251 field, model estimates alone however cannot capture the spatial and temporal variability outside
252 of quiet-time and at the scale that animals moving near the Earth's surface might experience the
253 geomagnetic field.

254 **2. Data and Methods**

255 MagGeo is an open-source tool that combines model estimates and satellite data and
256 attaches it to wildlife tracking data anywhere on the Earth's surface from November 2013 to
257 present. Model estimates are available for any 3-D location and timestamp of an animal tracking
258 fix. MagGeo uses the CHAOS-7 model to estimate the core, crustal and magnetosphere
259 contributions of the geomagnetic field (Figure 1B). CHAOS-7 model estimates do not calculate
260 ionospheric contributions. There are other models, such as the Swarm Comprehensive models
261 (Sabaka et al., 2020), which provide estimates for the ionosphere. The flexibility of the MagGeo
262 framework can allow for replacing the CHAOS-7 values with other data sources that provide
263 complete model estimates of the geomagnetic field based on user preference and research
264 objectives. Model estimates however will include only averaged quiet time values for the
265 ionosphere and subsequently will not capture the local, real-time variation. We include the Swarm
266 satellite data to introduce this local and temporal variability to CHAOS-7 model estimates.

267 For the CHAOS-7 model, the contributions from the core, lithosphere and magnetosphere
268 are added to create an estimate of geomagnetic values at the ground level:

$$269 \quad CH_g = CH_g^C + CH_g^L + CH_g^M \quad [1]$$

270 Where CH represents the CHAOS-7 model estimates, the subscript g represents geomagnetic
271 estimates at ground level altitude and the superscripts represent the different geomagnetic source
272 components (C = core, L = lithosphere, M = magnetosphere). As the model estimates are designed
273 to be continuous, they do not require spatial interpolation.

274 Raw geomagnetic data collected by satellites from ESA's Swarm constellation (European
275 Space Agency, 2021; Friis-Christensen et al., 2006) are also freely available. It is unlikely however

276 that geomagnetic values at satellite altitude will represent the geomagnetic field as experienced by
 277 an animal at the ground level. To correct for this altitude difference, we calculate satellite residuals
 278 by subtracting CHAOS-7 model estimates at satellite altitude for the core, lithosphere, and
 279 magnetosphere contributions.

$$280 \quad SW_s^{Res} = SW_s - CH_s^C - CH_s^L - CH_s^M \quad [2]$$

281 Where SW represents raw geomagnetic data collected by Swarm satellites, the subscript s
 282 represents values collected or estimated at satellite altitude and the superscripts represent the
 283 different geomagnetic source components (C, L, or M) or the satellite residuals (Res). The Swarm
 284 satellite residuals (SW_s^{Res}) primarily represent ionosphere contributions at satellite altitude though
 285 they are ultimately a combination of ionosphere, magnetosphere, crust, and other smaller
 286 influences on the geomagnetic field. SW_s^{Res} introduces temporal variability with fine resolution to
 287 capture the dynamic nature of the geomagnetic field outside of quiet-time values as estimated by
 288 geomagnetic models.

289 Given the satellite orbit however, it is unlikely that the satellite will be directly above a
 290 location on Earth for a specific timestamp. Therefore, geomagnetic data collected by satellites
 291 require interpolation to attach to an animal tracking fix. SW_s^{Res} can be interpolated to the animal
 292 tracking point by creating a space-time kernel. This kernel is a space-time cylinder where the radius
 293 of the cylinder has spatial dimensions, and the height has a temporal dimension (Figure 3A). Based
 294 on the Swarm satellites' polar orbits, the kernel's spatial boundary (the size of the cylinder's base,
 295 R in Figure 3A) varies with latitude, with smaller spatial boundaries at higher latitudes
 296 (approximately 900 km) compared to equatorial latitudes (approximately 1800 km). The kernel's
 297 temporal boundary (the height of the cylinder) is +/- 4 hours (ΔT) from the tracking fix, again

298 based on the properties of the polar orbit and to ensure that sufficient satellite data are present at
299 lower latitudes (Benitez-Paez et al., 2021).

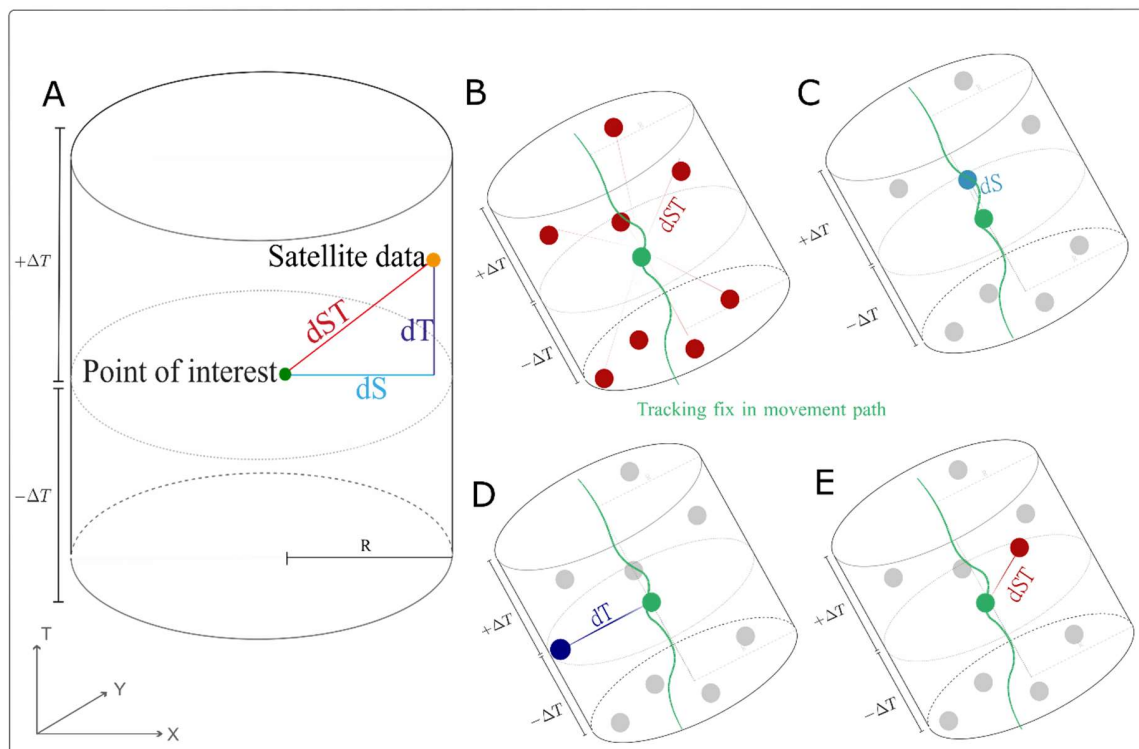
300 SW_s^{Res} within the space-time kernel are then linked to the animal tracking fix using a
301 spatiotemporal interpolation method (Figure 3B). Benitez-Paez et al. (2021) proposed inverse
302 distance weighting (IDW) where the space-time distance (dST) is calculated to account for both
303 the distance in space (measured in km; dS) and time (measured in seconds; dT) between the
304 satellite residual and the tracking fix. Data points closest in space-time distance (lowest dST) are
305 weighted higher than those farther away and the sum of all weights in each spacetime kernel is 1.
306 We propose three alternative nearest neighbour methods that are both simpler, and potentially
307 more accurate for spatiotemporal interpolation of satellite residuals to wildlife tracking fixes.

308 The nearest neighbour in space (NNS) interpolation method uses the residuals from the
309 satellite data point closest in space (lowest dS within the space-time kernel) (Figure 3C). It follows
310 that to create the nearest neighbour in time (NNT) and space-time (NNST) interpolation, we use
311 residuals of the satellite point closest in time and spacetime to the point of interest (lowest dT and
312 dST within the space-time kernel respectively; Figure 3D- 3E). Finally, after interpolation to the
313 animal tracking fix, we then add the satellite residuals with the CHAOS-7 model estimates at
314 ground altitude for the core, lithosphere, and magnetosphere contributions to get the final MagGeo
315 output.

$$316 \quad MG = SW_s^{Res} + CH_g \quad [3]$$

317 Thus, to create a complete model of the geomagnetic field for a 3-D location and timestamp of
318 an animal tracking fix which includes core, crustal, ionospheric and magnetospheric
319 contributions, MagGeo combines model estimates and satellite residuals. For simplicity, we will
320 refer to the fused model estimate and satellite residual outputs as MG and the CHAOS-7 model

321 estimates at ground altitude as CH for the remainder of the paper. We can use MG and CH
 322 values measured in the North-East-Centre (NEC) coordinate system to calculate the four
 323 components of the geomagnetic field that are relevant for animal migration (FHID). We perform
 324 our error analysis on these components as their values are more applicable to MagGeo users
 325 whose primary objective will likely focus on movement ecology research questions. For a similar
 326 geophysical-centered error analysis on the XYZ orthogonal components of geomagnetic model
 327 estimates, readers can refer to Beggan (2022).



328

329 **Figure 3.** Current and modified spatiotemporal parameters of the spacetime cylinder used by MagGeo. **A.**
 330 The space-time cylinder calculating the distance in space (dS ; light blue), time (dT ; dark blue) and
 331 spacetime (dST ; red) between point of interest (green) and satellite point (orange). X and Y axis represent
 332 spatial dimensions whereas the Z-axis represents temporal dimension. Figures **B-D** represent the four
 333 spatiotemporal interpolation methods used to attach satellite geomagnetic data (residuals) to an animal
 334 tracking fix (such as movement path of a migratory bird collected by a GPS tag). Figure **B** represents the
 335 concept of inverse-distance weighting and Figures **C-E** represent nearest neighbour iterations for satellite
 336 points closest to the animal tracking fix in space (**C**), time (**D**) and spacetime (**E**).

337 *2.1 Data Preparation*

338 Our analysis centers on the assumption that data from INTERMAGNET observatories
339 represent the best available measurement of the geomagnetic field at their location on the Earth's
340 surface (Beggan, 2022; Kerridge, 2001). Data from these terrestrial observatories are acutely
341 influenced by the local crustal field, which is not captured by model estimates or satellite data but
342 might be detected by animals moving at this local spatial scale. Additionally, INTERMAGNET
343 observatories collect high temporal resolution geomagnetic data and are rigorously calibrated
344 (Kerridge, 2001). We compare geomagnetic values from INTERMAGNET observatories (OBS)
345 against MG outputs interpolated to the station location for the same timestamp. The objective is to
346 compare MG and OBS values for the four geomagnetic components (F, H, D, I) to assess MagGeo
347 accuracy at ground altitude. Our analysis is conceptually consistent with geophysical studies that
348 test, calibrate and validate satellite (Beggan et al., 2013; Macmillan & Olsen, 2013; Ridley &
349 Macmillan, 2014) and model data (Chulliat et al., 2015; Finlay et al., 2020).

350 We acquired minute-mean observatory geomagnetic data for all available stations for seven
351 years (2014-2020). We compiled a dataset to test MagGeo under the full range of geomagnetic
352 activity levels (i.e., Kp 0 – 9). We term this dataset “All Kp.” We included data from 60 days a
353 year uniformly sampled across all twelve months for seven years resulting in a total of 420 days
354 of data. We obtained geomagnetic data at three time points each day equally spaced 8 hours apart
355 (Table 1). We had fewer INTERMAGNET stations for later years as there is usually a delay
356 between station measurements and access to the final geomagnetic dataset.

357 To test MagGeo specifically during periods of high geomagnetic activity, we compiled a
358 “High Kp” dataset. To build the High Kp dataset, we acquired data for all days in 2014-2020 with
359 high geomagnetic activity (Kp > 6 for 6 or more hours) (Space Weather Live, 2021). We subset

360 this dataset to include only data where the satellite recorded $K_p > 6$ to further filter out quieter
 361 periods even during a day classified as having overall high geomagnetic activity. The High K_p
 362 dataset ($n = 393,054$) was substantially smaller than the All K_p dataset ($n = 6,327,537$). There
 363 were fewer days with High K_p in 2014 and in 2018 to 2020. These periods were less
 364 geomagnetically active as they were in the quieter part of previous solar cycle (Kakad et al., 2020).

365 **Table 1.** Datasets used for MagGeo error and accuracy analysis. All data are minute-mean for
 366 2014 to 2020. All K_p includes data from all K_p levels whereas High K_p includes only data from
 367 geomagnetically active periods ($K_p > 6$)

Year	All K_p	High K_p	INTERMAGNET stations
2014	899 967	8 602	109
2015	950 346	206 448	111
2016	1 044 477	78 467	103
2017	1 023 750	65 259	100
2018	892 808	26 363	95
2019	888 681	7 915	97
2020	627 508	8 602	88
<i>n total</i>	<i>6,327,537</i>	<i>393,054</i>	

372

373 2.2 Accuracy Assessment

374 To test the performance of the different spatiotemporal interpolation strategies and data
 375 structures relative to one another we used two accuracy measures. The first is the absolute
 376 difference (d) between the MagGeo output (X_{mg}) and the corresponding observatory value (X_{obs})
 377 for each timestamp:

$$378 \quad d = |X_{mg} - X_{obs}| \quad [4]$$

379 Lower values of absolute error correspond to better agreement between the MagGeo output and
 380 the INTERMAGNET data.

381 The second measure (alpha; α) is the absolute difference between the standardized MagGeo
 382 output and the standardized observatory output for each timestamp (Ridley & Macmillan, 2014):

$$383 \quad \alpha = \left| \left(\frac{X_{mg} - \bar{X}_{mg}}{\sigma_{mg}} \right) - \left(\frac{X_{obs} - \bar{X}_{obs}}{\sigma_{obs}} \right) \right| \quad [5]$$

384 Where \bar{X}_{mg} and σ_{mg} are the mean and standard deviation respectively for the interpolated
 385 MagGeo values at a station while \bar{X}_{obs} and σ_{obs} are the mean and standard deviation respectively
 386 of the INTERMAGNET values at the same station. As with the absolute difference measure, lower
 387 alpha values correspond to better agreement between the MagGeo output and observatory values.
 388 The alpha measure is useful for identifying how well MagGeo captures relative patterns in the
 389 geomagnetic data instead of just the absolute difference. For example, during a geomagnetic storm,
 390 both X_{mg} and X_{obs} are very different from their respective means (\bar{X}_{mg} and \bar{X}_{obs}). While the
 391 absolute difference between X_{mg} and X_{obs} might be large during these storms, if the sudden change
 392 in geomagnetic value is captured by both sources, the α value will be low thus making it possible
 393 for MagGeo outputs and observatory outputs to have high absolute difference error but low alpha
 394 values. In this case, the pattern would suggest that MagGeo is able to capture the temporally
 395 dynamic nature of the geomagnetic field like the observatory data irrespective of any consistent
 396 offsets between the two sources.

397 We removed data from 8 stations (Figure S1) because their absolute difference error was
 398 consistently greater than 2 standard deviations (95% quantile) for any three of the orthogonal
 399 components (eg., DIF, XYZ or DHZ) for more than 6 months' worth of data (Beggan, 2022). We
 400 also removed data from 3 days (2017-09-08, 2018-08-26, and 2018-08-27) which had high daily
 401 error across all stations reflecting the impact of very strong geomagnetic storms on these days.

402 We calculated the error measures (d and α) for MagGeo values for each spatiotemporal
403 interpolation method (IDW, NNT, NNS, and NNST) and underlying data structure (CH and MG).
404 We compare across interpolation methods and data structure using summary statistics, but also by
405 recording the proportion of data records where each interpolation method and data structure had
406 the lowest error values (“best performance”). We calculated Pearson’s correlation coefficients to
407 compare the best performing MagGeo algorithm with observatory values overall at each station
408 for each geomagnetic component.

409 We fit generalized linear mixed-effect models (GLMMs) with the dependent variable as
410 either error metric (d and α) and by using data values from the best performing interpolation
411 method and data structure combination. To reduce temporal autocorrelation, we created hourly
412 averages from our minute-mean data resulting in a total of 107,116 data points. For fixed effects,
413 we included three variables that account for geomagnetic field behaviour: Kp, time of day, and
414 latitude. For simplicity, we categorized time-of-day into two categories: day as 7:00AM to 7:00PM
415 local time and night as 7:00PM to 7:00AM local time to allow sufficient variation in sunset and
416 sunrise times for stations differing by latitude. Solar wind influences the geomagnetic field activity
417 and is reflected as a high Kp value which is more likely during the day and at polar latitudes
418 (Campbell, 2003; Hulot et al., 2010; Lanza & Meloni, 2006). We also included two additional
419 fixed effects in our model that address how MagGeo space-time kernel parameters may influence
420 error: the geographical distance between the INTERMAGNET station and the satellite data point
421 (km) and the temporal difference (minutes) between the timestamp at the INTERMAGNET station
422 data and the nearest Swarm satellite pass.

423 We used station ID ($n=109$) as a random effect for all models as it is likely that values from
424 individual ground stations are heavily influenced by local crustal field conditions (Beggan, 2022;

425 Lesur et al., 2016). Additionally, each station has subtle differences in collection and reporting of
426 geomagnetic data (St-Louis, 2012). We tested all possible combinations of the fixed effects and
427 chose the best model for each geomagnetic component based on the lowest Akaike information
428 criterion (AIC) values. In addition to the coefficient values (β), standard error (SE) and p-values,
429 we also calculated the marginal and conditional R^2 values for the best performing models. Marginal
430 (R^2_m) is the proportion of the variance explained by the fixed effects, and the conditional R^2 (R^2_c)
431 is the overall proportion of the variance explained by both the fixed and random effects (Nakagawa
432 & Schielzeth, 2013).

433 To demonstrate the difference between the geomagnetic data sources for studying long-
434 distance animal migration, we used GPS tracking data from one White Stork *Ciconia ciconia*
435 individual from the 2017 spring migration period (Carlson et al., 2021). We resampled the tracking
436 data to hourly intervals when the bird was in flight (speed > 5 km/h) as this state likely reflects
437 when birds are using geomagnetic field values to make movement decisions (Acácio et al., 2022;
438 Chernetsov, 2017). We attached geomagnetic values from the nearest INTERMAGNET station to
439 the bird's location. We compared these observatory values with the MagGeo outputs from the best
440 performing interpolation method and data structure for the same location.

441 *2.3 Tools and data availability*

442 MagGeo is available as a GitHub repository (Benitez & Long, 2022). For our analysis, we
443 modified scripts from MagGeo 1.0 (Feb 2021). MagGeo uses two Python packages for
444 geomagnetic data acquisition. The ESA-VirES Client package connects to the VirES servers to
445 acquire satellite residuals (Smith, 2020) whereas the chaomagpy package accesses the CHAOS-
446 7 model estimates through the VirES server (Kloss, 2021). We used the Swarm Magnetic Earth
447 Jupyter notebooks to fetch geomagnetic data from ground observatories which we accessed via the

448 British Geological Survey FTP server, though a VirES-based access method is currently available
449 as well (https://github.com/Swarm-DISC/Swarm_notebooks). To fit our general linear mixed-
450 effect models, we used the “lme4” R package (Bates, 2010). Finally, we used the “dredge” function
451 from the “MuMIn” R package to test all possible combinations of fixed effects (Bartoń, 2022).
452 We accessed White Stork GPS data from Carlson et al. (2021) which are available on Movebank
453 (Kays et al., 2022).

454 **3. Results**

455 For each of the four spatiotemporal interpolation methods, there was little variation in
456 the median, mean, standard deviation, or skew across all geomagnetic components, for both
457 accuracy measures and during variable (All Kp) and high geomagnetic activity (High Kp)
458 (Table 2). The variation in mean absolute difference between the four interpolation methods
459 was within 10 nT for scalar intensity geomagnetic components (F and H) and within 1 degree
460 for angular directional geomagnetic components (D and I) (Table 2). Furthermore, almost all
461 categories showed positive skew (with median < mean), suggesting that mean values may be
462 influenced by a few data points with unusually high error. For both scalar components, while
463 values between the interpolation methods were similar, NNT often had the highest median and
464 mean values. Across each data record (unique location and timestamp) however, the nearest
465 neighbour in time method always had the highest occurrence (%) of lowest error (“best
466 performance”) for all components during periods of variable and high geomagnetic activity
467 (Table 3). This distinction was more evident during variable geomagnetic activity where NNT
468 had the lowest error about 40% of the time among the four interpolation methods compared to
469 periods of high geomagnetic activity where NNT on average had the lowest error 30% of the
470 time (Table 3). Thus, all interpolation methods had similar central tendencies (Table 2), but
471 NNT consistently had the best performance (Table 3). Therefore, we used the NNT
472 interpolation method to subsequently test the difference between CH and MG

473 When separated by station and arranged by latitude, we found that certain stations have
474 greater variation in error than others (large interquartile range for individual station box plot)
475 (Figure 4 and S2). In general, stations at higher latitudes have a greater variation in differences
476 compared to stations closer to the equator. We also observe that these error patterns are
477 consistent between the four interpolation methods such that if a station has high error variability

478 for total intensity, this pattern will be replicated across all interpolation methods (Figure 4 and
479 S2).

480 Between the two data structures, there was little variation in the median, mean, standard
481 deviation, or skew across all geomagnetic components, for both accuracy measures and during
482 variable (All Kp) and high geomagnetic activity (High Kp) (Table 4). With a few exceptions,
483 difference between CH and MG for the mean and median absolute difference error was
484 approximately within 10 nT for the scalar components and within 1 degree for the angular
485 components (Table 4). For the absolute difference error metric, apart from horizontal intensity,
486 CH had equal or lower median and mean error than MG (Table 4). During both variable and
487 high geomagnetic activity however, MG had either equal or lower mean and median alpha
488 values. Positive skew during variable geomagnetic conditions is also higher for MG alpha
489 values suggesting that the reported mean is being skewed by a few instances of very high error
490 (Table 4).

491 For 11 out of 16 categories based on geomagnetic components and activity, MG has slightly
492 better performance than CH (Table 5). However, there is little difference in performance across
493 all geomagnetic components, activity and accuracy measures since the overall average
494 performance is 48.3% for CH and consequently 51.7% for MG. The average alpha error per
495 station is also around 1 unit during variable geomagnetic activity and slightly higher during
496 high geomagnetic activity (Figure 5 and Figure S3). Except for declination, MG has lower
497 station-wide alpha error than CH (Figure 5 and Figure S3). Additionally, there is a log-linear
498 increase in Swarm satellite contribution to the MG data (e.g., increasing residual values)
499 associated with an increase in geomagnetic activity (Figure S4).

500 When using the best performing MagGeo algorithm (NNT interpolation and MG data
501 structure), MagGeo values for all geomagnetic components were overall highly correlated with

502 observatory values overall ($r > 0.9$ for FHID). Out of 109 stations and for all geomagnetic
503 components, at least 50% of all stations had correlation values greater than 0.5 though the
504 percentage of stations varied by geomagnetic component (F: 88%, H: 53%, D: 78% and I: 66%
505 of stations had correlation values greater than 0.5) (Figure S5). Horizontal intensity and
506 inclination had the lowest number of stations with high correlations though stations with low
507 correlations were typically at higher latitudes.

508 **Table 2.** Median, mean, standard deviation and skew of error for current and modified MagGeo spatiotemporal interpolation methods: IDW, NNS,
 509 NNST and NNT. Results are for all geomagnetic components (FHID) with minute-mean data from 2014-2020. “All Kp” includes data from all
 510 Kp levels whereas “High Kp” includes only data from geomagnetically active periods (Kp>6). Results are presented for both accuracy measures:
 511 absolute difference (d) and alpha (α).

Component	d								α							
	All Kp				High Kp				All Kp				High Kp			
	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew
Total intensity																
IDW	79.2	127.4	135.4	2.1	101.5	258.8	592.3	6.0	0.5	0.7	0.8	2.4	0.8	1.0	0.8	2.6
NNS	79.4	127.4	135.0	2.1	101.6	259.7	593.2	5.9	0.5	0.7	0.8	2.3	0.8	1.0	0.9	5.2
NNST	79.4	127.6	135.3	2.1	100.8	258.6	592.5	5.9	0.5	0.7	0.8	2.4	0.8	1.0	0.9	4.2
NNT	79.7	128.0	135.7	2.1	108.4	267.3	591.3	5.9	0.5	0.8	0.8	2.7	0.8	1.0	0.9	4.0
Horizontal intensity																
IDW	211.3	211.2	133.1	1.0	284.4	371.1	393.2	4.1	0.6	0.9	0.9	2.8	0.9	1.1	0.9	2.0
NNS	211.5	211.7	134.0	1.1	284.3	374.1	401.3	4.0	0.6	0.9	0.9	2.9	0.8	1.0	0.9	4.5
NNST	211.7	212.0	134.2	1.1	285.5	375.6	401.7	4.0	0.6	0.9	0.9	2.9	0.9	1.0	0.9	3.6
NNT	211.5	211.9	133.4	1.0	298.0	382.8	392.6	4.0	0.6	0.9	0.9	2.9	0.9	1.1	0.9	3.6
Inclination																
IDW	0.3	0.3	0.2	2.6	0.4	0.5	0.5	4.0	0.6	0.8	0.9	2.9	0.9	1.1	0.9	2.0
NNS	0.3	0.3	0.2	2.6	0.4	0.5	0.5	4.0	0.6	0.8	0.9	3.0	0.8	1.0	0.9	4.7
NNST	0.3	0.3	0.2	2.6	0.4	0.5	0.5	4.0	0.6	0.8	0.9	3.0	0.8	1.0	0.9	3.6
NNT	0.3	0.3	0.2	2.6	0.4	0.5	0.5	4.0	0.6	0.8	0.9	2.9	0.9	1.1	0.9	3.6
Declination																
IDW	0.1	0.4	0.6	3.8	0.3	0.9	2.4	13.2	0.7	0.9	0.8	2.2	0.9	1.0	0.9	2.3
NNS	0.2	0.4	0.7	3.6	0.3	0.9	2.4	12.1	0.7	0.9	0.9	2.3	0.8	1.0	0.9	5.4
NNST	0.2	0.4	0.7	3.6	0.3	0.9	2.4	12.2	0.6	0.9	0.8	2.4	0.8	1.0	0.9	4.1
NNT	0.2	0.4	0.7	3.7	0.3	1.0	2.5	10.9	0.7	0.9	0.9	2.7	0.8	1.0	0.9	4.2

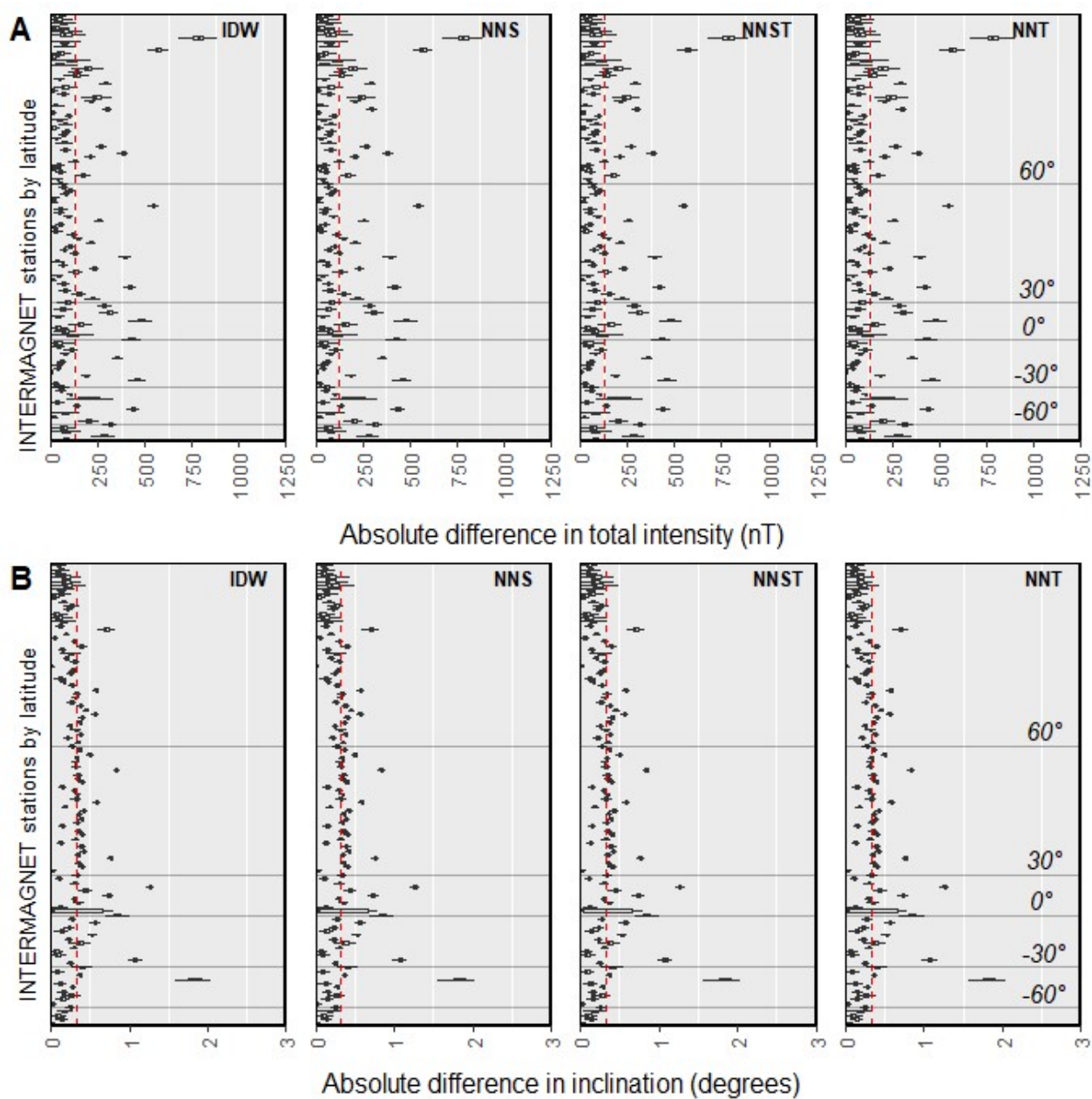
512

513 **Table 3.** Percent of lowest error between four MagGeo spatiotemporal interpolation methods:
 514 IDW, NNS, NSST, and NNT. Bolded rows represent the interpolation method that had the overall
 515 highest frequency of lowest error (best performance). Results shown for all geomagnetic
 516 components (FHID) with minute-mean data from 2014-2020. “All Kp” includes data from all Kp
 517 levels whereas “High Kp” includes only data from geomagnetically active periods (Kp>6). Results
 518 are for both accuracy measures: absolute difference (d) and alpha (α).

519

<i>Component</i>	Percent (%) with lowest error			
	All Kp		High Kp	
	d	α	d	α
Total Intensity				
IDW	17.8	18.5	22.0	27.8
NNS	31.4	19.3	33.4	22.7
NNST	8.4	22.5	5.6	22.4
NNT	42.4	39.7	39.1	27.2
Horizontal intensity				
IDW	19.4	18.6	21.5	22.6
NNS	30.8	20.4	34.2	24.8
NNST	8.9	19.9	5.3	22.4
NNT	40.9	41	39.1	30.2
Inclination				
IDW	19	19.5	20.9	23.2
NNS	30.6	19.9	34.1	24.5
NNST	8.7	21.1	5.0	22.9
NNT	41.7	39.4	39.9	29.5
Declination				
IDW	21.3	20.3	26.8	25.8
NNS	30.5	19.5	32.9	24.6
NNST	8.2	20.7	5.1	21.7
NNT	40.1	39.5	35.2	27.9

520



521

522 **Figure 4.** Absolute difference metric (d) for each individual INTERMAGNET station arranged
 523 by latitude from northernmost (top) to southernmost (bottom). Each panel represents one of the
 524 four possible MagGeo spatiotemporal interpolation methods: IDW, NNS, NNST and NNT.
 525 Dotted red lines represent overall average absolute difference. Figures are arranged by
 526 geomagnetic component: total intensity (A) and inclination (B). For similar figures for horizontal
 527 intensity and declination, see Figure S2.

528 **Table 4.** Median, mean, standard deviation and skew of error for current and modified MagGeo data structures: CH (CHAOS-7
 529 estimates) and MG (CHAOS-7 estimates and Swarm residuals). Results are for all geomagnetic components (FHID) with minute-mean
 530 data from 2014-2020. “All Kp” includes data from all Kp levels whereas “High Kp” includes only data from geomagnetically active
 531 periods (Kp>6). Results are for both accuracy measures: absolute difference (d) and alpha (α).

Component	d								α							
	All Kp				High Kp				All Kp				High Kp			
	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew	Median	Mean	\pm SD	Skew
Total intensity																
CH	78.6	126.9	135.3	2.1	105.0	261.8	595.3	5.9	0.6	0.8	0.8	2.1	1.0	1.1	0.8	0.9
MG	79.7	128.2	153.8	65.8	108.4	267.3	591.3	5.9	0.5	0.8	0.2	3.7	1.0	1.1	0.8	1.0
Horizontal intensity																
CH	212.0	211.1	131.7	1.0	297.8	365.9	379.6	4.5	0.8	1.0	0.8	2.1	1.0	1.1	0.9	1.0
MG	211.7	212.4	134.9	2.1	298.0	382.8	392.6	4.0	0.6	0.9	1.0	3.7	0.9	1.1	0.9	1.3
Inclination																
CH	0.3	0.3	0.2	2.6	0.4	0.5	0.5	4.1	0.7	0.9	0.8	2.2	1.0	1.1	0.8	1.0
MG	0.3	0.3	0.2	3.7	0.4	0.5	0.5	4.0	0.6	0.8	0.9	3.7	0.9	1.1	0.9	1.2
Declination																
CH	0.1	0.3	0.6	3.9	0.2	0.8	2.3	13.8	0.7	0.9	0.8	1.8	1.0	1.1	0.8	1.0
MG	0.2	0.4	0.7	4.1	0.3	1.0	2.5	10.9	0.7	0.9	0.9	3.3	0.9	1.1	0.9	1.2

532

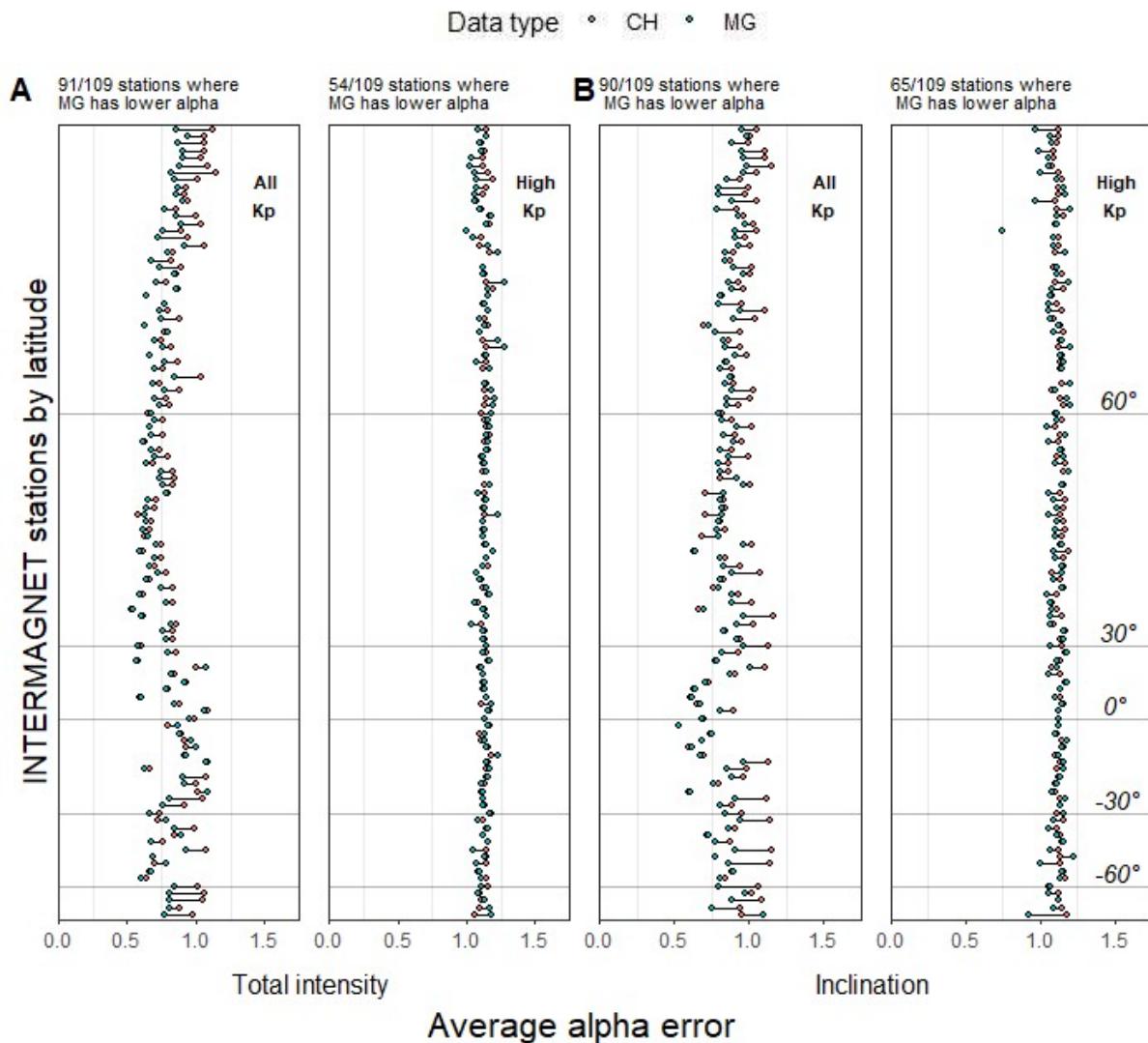
533 **Table 5.** Percent of lowest error between two MagGeo data structures: CH (CHAOS-7
 534 estimates) and MG (CHAOS-7 estimates and Swarm residuals). Bolded rows represent the
 535 data type that had the highest percentage of lowest error for each category (best performance).
 536 Results are for all geomagnetic components (FHID) with minute-mean data from 2014-2020.
 537 “All Kp” includes data from all Kp levels whereas “High Kp” includes only data from
 538 geomagnetically active periods (Kp>6). Results are for both accuracy measures: absolute
 539 difference (d) and alpha (α).

<i>Component</i>	Percent (%) with lowest error			
	All Kp		High Kp	
	d	α	d	α
Total Intensity				
CH	52.8	41	50.8	48.9
MG	47.2	59	49.2	51.1
Horizontal intensity				
CH	47.9	40.8	47.6	48.6
MG	52.1	59.2	52.4	51.4
Inclination				
CH	50.7	41.4	47.3	48.1
MG	49.3	58.6	52.7	51.9
Declination				
CH	55.6	48	58.2	45.3
MG	44.4	52	41.8	54.7

540

541

542



543

544 **Figure 5.** Comparison of average alpha measure (α) between CH (red dots) and MG (blue
 545 dots) for each individual INTERMAGNET station arranged by latitude from northernmost
 546 (top) to southernmost (bottom). “All Kp” includes data from all Kp levels whereas “High
 547 Kp” includes only data from geomagnetically active periods ($Kp > 6$). Individual plot titles
 548 indicate count of stations where MG had lower alpha error. Figures are arranged by
 549 geomagnetic component: total intensity (**A**) and inclination (**B**). For similar figures for
 550 horizontal component of intensity and declination, see Figure S3.

551 *3.1 Factors influencing error and accuracy structure*

552 For distance in space, most satellite data points are between 1000-1250 km away from the
553 INTERMAGNET station (Figure 6). This pattern might be a result of the clustering of stations at
554 mid-latitudes in Europe (Figure 2 and 6B) who will have similar space-time kernel parameters and
555 subsequently error. The low hourly error at smaller distances might reflect that MagGeo can
556 accurately capture geomagnetic patterns if the satellites are close in space to the INTERMAGNET
557 station (Figure 6A). Conversely, the low hourly error at high distances might be indicative of
558 stations near the equator who have larger space-time kernels but are also found at latitudes where
559 there is lower geomagnetic activity (Figure 6B).

560 There is little overall variation in the distance in time between the INTERMAGNET station
561 and satellite data point though there are two peaks at 0-30 minutes and 190-220 minutes (Figure
562 7). This clustering is likely related to the +/- 4-hour parameter of the space-time kernel where the
563 nearest satellite data point is either directly above the INTERMAGNET station (0-30 minutes) or
564 will just meet the +/-4-hour cut-off by either passing over the INTERMAGNET station 4 hours
565 ago or passing over it 4 hours later (190-220 minutes).

566 Random effects (individual stations) and fixed effects (geomagnetic activity and time of
567 day) together explain most of the absolute difference between geomagnetic values collected at
568 INTERMAGNET stations and outputs from MagGeo (conditional R^2 is close to 0.9-1.0 for all
569 geomagnetic components, Table 6). Apart from inclination (marginal R^2 is 0.5), variation at
570 individual stations (random effects) explains most of the difference (marginal R^2 is equal to or
571 below 0.1). Generally, INTERMAGNET data from stations at higher absolute latitudes are
572 consistently different from MagGeo outputs, especially for the angular directional geomagnetic
573 components such as inclination and declination (Table 6, Figure 4B).

574 Random and fixed effects do not explain most of the variation in the alpha error
575 (conditional R^2 is equal to 0.11, Table 7). For these models however, fixed effects explain all the
576 difference between MagGeo outputs and INTERMAGNET stations (marginal R^2 equal to
577 conditional R^2 , Table 7). For example, for all geomagnetic components, high geomagnetic activity
578 leads to higher error (Table 7). These models also suggest that alpha error is higher during the day
579 compared to nighttime. None of the best models (lowest AIC) for any geomagnetic component
580 with absolute difference as a dependent variable have distance in space as a fixed effect (blank
581 rows in Table 6) whereas only total intensity and inclination have distance in time included as a
582 fixed effect. This is in contrast with the best models with alpha error as a dependent variable since
583 almost all components have both distance in space and time as fixed effects in their models. While
584 not always statistically significant, in models where these variables are included distance in space
585 and time between satellite pass and INTERMAGNET station have an impact on error such that an
586 increase in distance leads to larger alpha error.

587 **Table 6.** Results of the generalized linear mixed-effect models of absolute difference as dependent variable, individual station as a
 588 random effect and fixed effects as Kp, absolute value of latitude, time of day and distance in space and time between satellite and point
 589 of interest. R^2_m and R^2_c refer to marginal and conditional R^2 respectively. A new model was fit for each geomagnetic component
 590 (FHID). Models are fit with combined model estimate and satellite residual data and NNT interpolation during periods of variable
 591 geomagnetic activity (All Kp). Minute-mean data are averaged into hourly values, and empty rows indicate that the fixed effect did not
 592 contribute to the final model (based on lowest AIC) for that geomagnetic component.

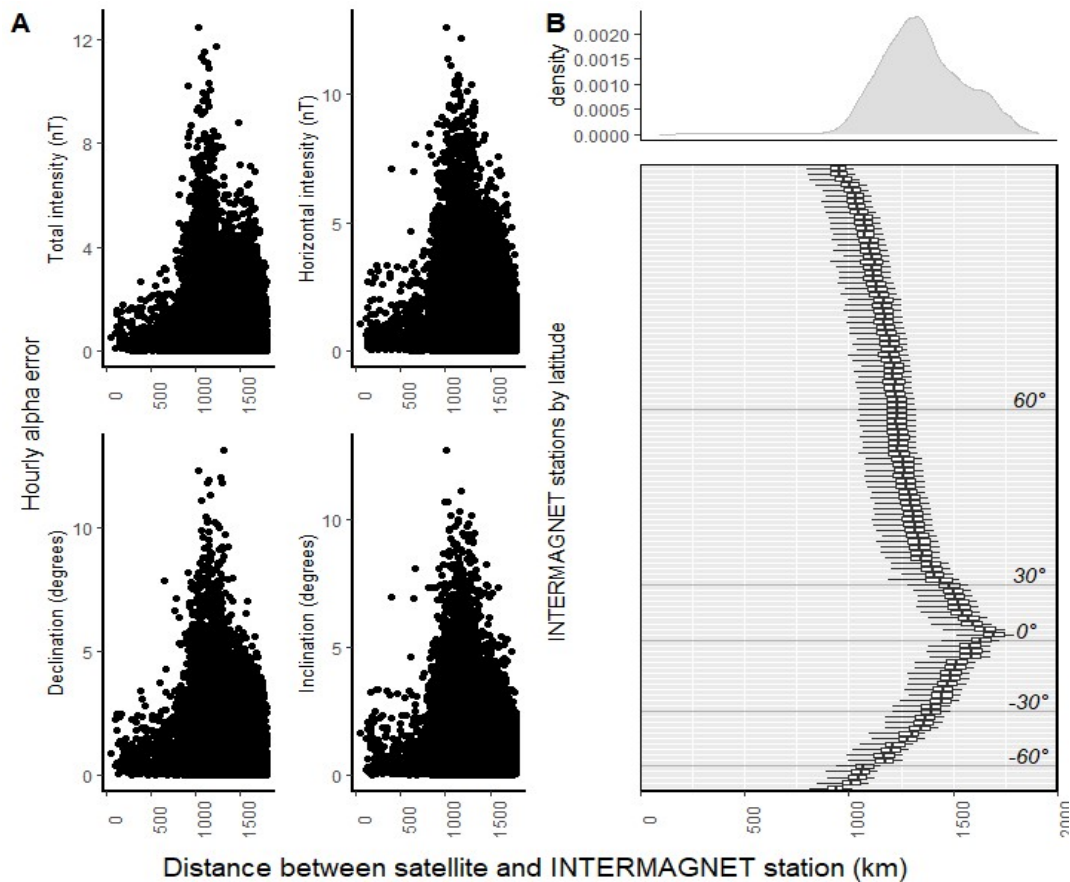
	Total intensity			Horizontal intensity			Inclination			Declination		
	β	SE	p	β	SE	p	β	SE	p	β	SE	p
<i>Intercept</i>	71.3	36.3	0.05	245.6	32.6	<0.01	-33.0	1.3	<0.01	-0.2	0.1	0.2
<i>Kp</i>	1.6	0.1	<0.01	5.5	0.1	<0.01	0.007	1.3E-04	<0.01	0.02	5.9E-04	<0.01
<i>Absolute Latitude</i>	1.4	0.8	0.07	-1.1	0.7	0.1	0.7	0.01	<0.01	0.01	0.003	<0.01
<i>Time of day (Ref: Day)</i>												
<i>Night</i>	0.8	0.2	<0.01	4.7	0.3	<0.01	0.008	3.3E-04	<0.01	-0.004	0.001	0.01
<i>Distance (km)</i>												
<i>Time difference (min)</i>	-0.003	0.0	0.03				-3.8E-06	2.4E-06	0.1			
R^2_m		0.03			0.02			0.5			0.1	
R^2_c		1.0			0.9			1.0			0.9	

593

594 **Table 7.** Results of the generalized linear mixed-effect models of alpha measure as dependent variable, individual station as a random
 595 effect and fixed effects as Kp, absolute value of latitude, time of day and distance in space and time between satellite and point of
 596 interest. R^2_m and R^2_c refer to marginal and conditional R^2 respectively. A new model was fit for each geomagnetic component (FHID).
 597 Models are fit with combined model estimate and satellite residual data and NNT interpolation during periods of variable geomagnetic
 598 activity (All Kp). Minute-mean data are averaged into hourly values, and empty rows indicate that the fixed effect did not contribute to
 599 the final model (based on lowest AIC) for that geomagnetic component.

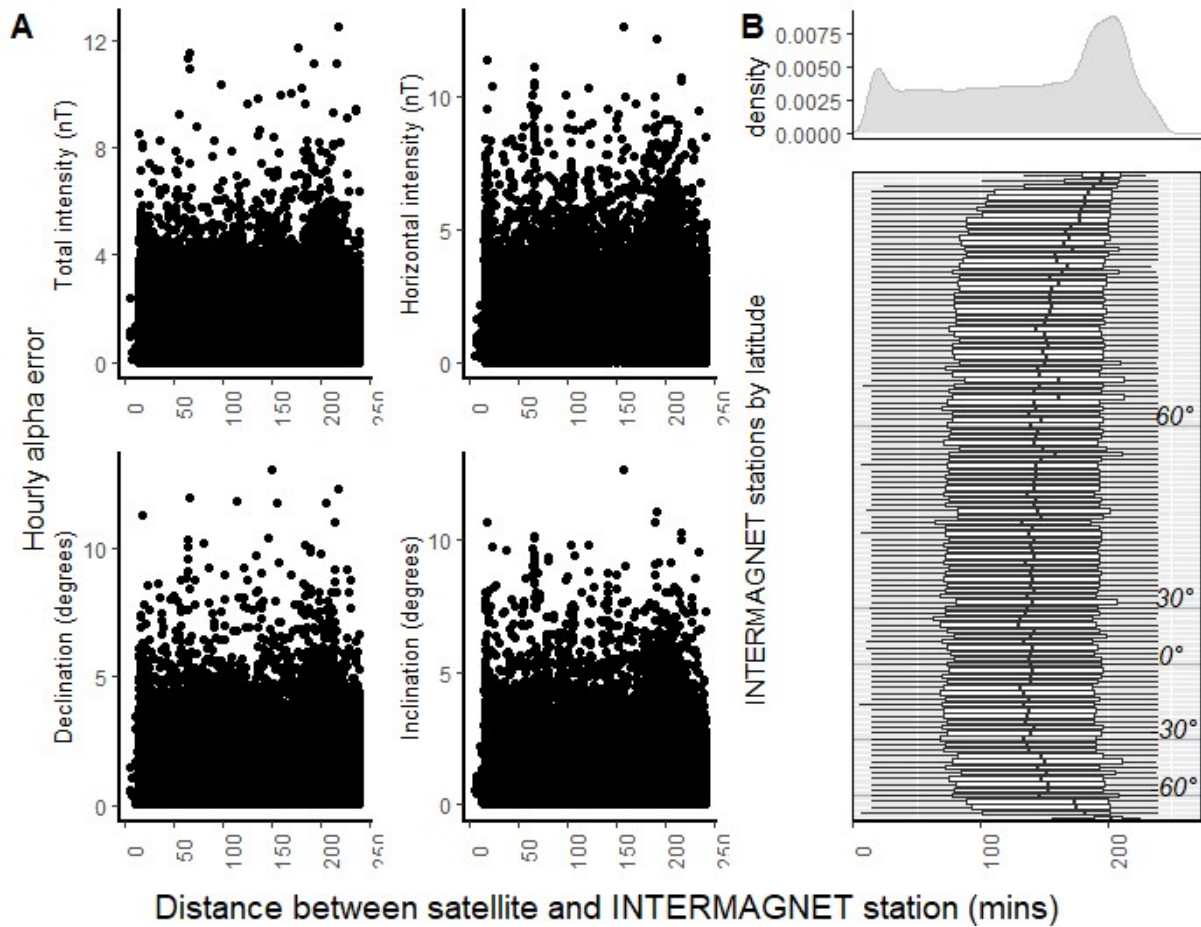
	Total intensity			Horizontal intensity			Inclination			Declination		
	β	SE	p	β	SE	p	β	SE	p	β	SE	p
<i>Intercept</i>	0.5	0.1	<0.01	0.5	0.03	<0.01	0.3	0.06	<0.01	0.8	0.01	<0.01
<i>Kp</i>	0.2	0.002	<0.01	0.2	0.002	<0.01	0.2	0.002	<0.01	0.1	0.002	<0.01
<i>Absolute Latitude</i>	-9.1E-04	6.4E-04	0.2				0.003	5.5E-04	<0.01			
<i>Time of day (Ref: Day)</i>												
<i>Night</i>	-0.1	0.005	<0.01	-0.2	0.005	<0.01	-0.1	0.005	<0.01	-0.3	0.01	<0.01
<i>Distance (km)</i>	4.4E-05	2.8E-05	0.1	4.2E-05	2.5E-05	0.1	4.4E-05	3.0E-05	0.1			
<i>Time difference (min)</i>	3.3E-04	3.4E-05	<0.01	1.8E-04	3.8E-05	<0.01	2.0E-04	3.6E-05	<0.01	2.5E-04	3.7E-05	<0.01
R^2_m		0.1			0.1			0.1			0.1	
R^2_c		0.1			0.1			0.1			0.1	

601



602

603 **Figure 6.** Distance in space between satellite pass and INTERMAGNET station (km) where (A)
 604 shows the distribution of hourly alpha error for each geomagnetic component (FHDI) and (B)
 605 shows the spread of the distance values for each individual station arranged by latitude from
 606 northernmost (top) to southernmost (bottom). Density plot in (B) mirrors the pattern seen in (A)
 607 where most distance values are between 1000-1500 km. Data are combined model estimate and
 608 satellite residual and interpolated to the station location using a NNT method during periods of
 609 variable geomagnetic activity (All Kp).

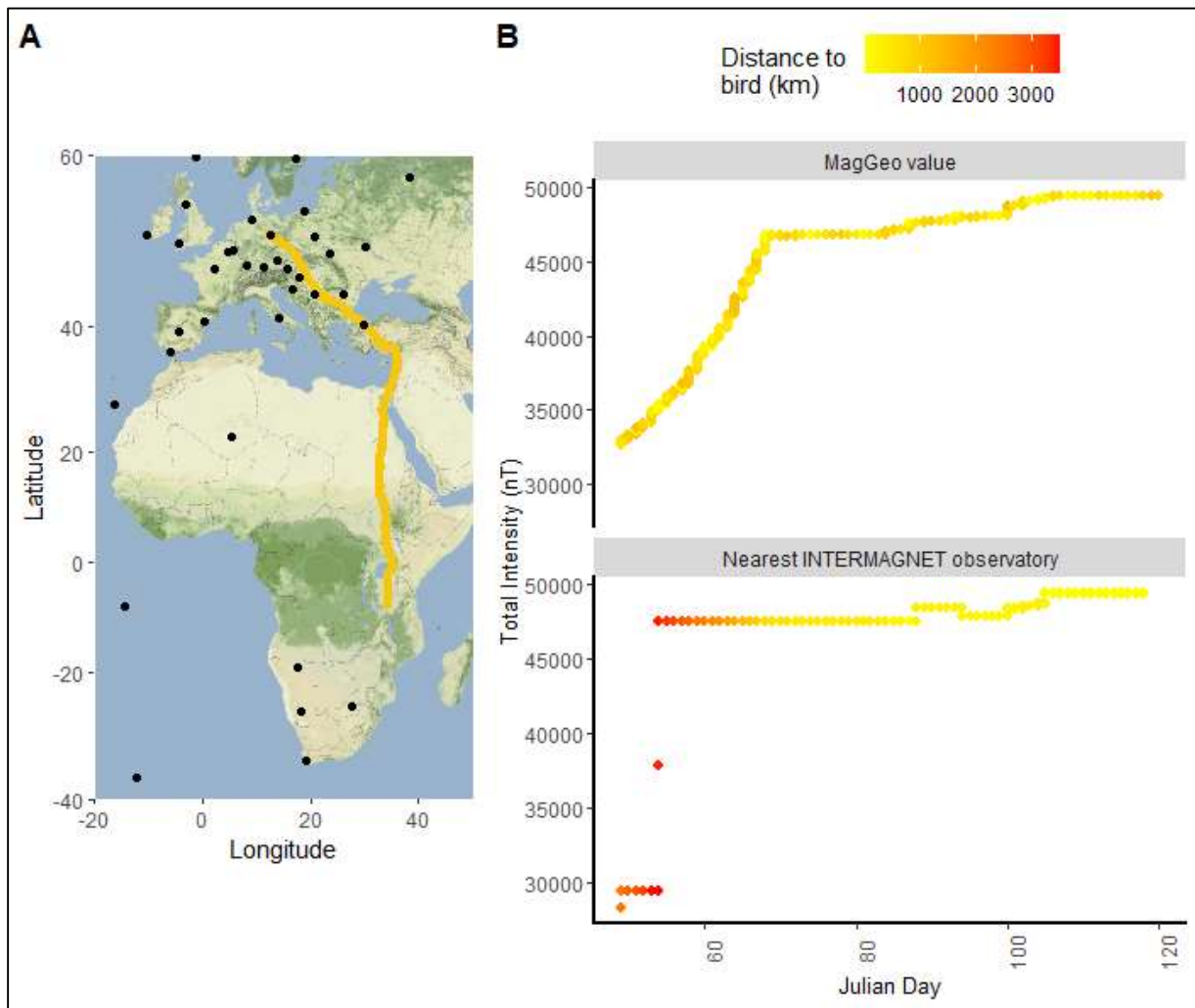


610

611 **Figure 7.** Time difference (mins) between satellite pass and INTERMAGNET station timestamp
 612 where (A) shows the distribution of hourly alpha error for each geomagnetic component (FHDI)
 613 and (B) shows the spread of time difference for each individual station arranged by latitude from
 614 northernmost (top) to southernmost (bottom). Data are combined model estimate and satellite
 615 residual and interpolated to the station location using a NNT method during periods of variable
 616 geomagnetic activity (All Kp).

617 *3.2 Attaching geomagnetic data to a migratory bird track*

618 We used movement data collected by a GPS tag carried by one White Stork individual
619 during spring migration as it moves from its wintering grounds in sub-Saharan Africa to its
620 breeding grounds in northeastern Europe by crossing the Sahara Desert. Due to the limited number
621 of stations in Africa, the distance between the bird's location and the nearest INTERMAGNET
622 station is high. This distance decreases as the bird nears its breeding grounds in Europe where
623 there is high station density. In this region, attaching data from the nearest INTERMAGNET
624 station to the bird's location effectively mirrors a continuous geomagnetic data surface. For
625 example, in the later stages of the bird's migration (after Julian Day 90), the curved lines at the
626 northern latitudes using observatory values appears like the curved lines using MagGeo outputs.
627 MagGeo outputs however use continuous model estimates across the world and can represent the
628 changes in geomagnetic values at every location as the bird moves across large distances. Even
629 when including the interpolated satellite residuals required to create the MG data framework, the
630 farthest distance between a satellite pass and a bird's location is still lower than the farthest distance
631 between an INTERMAGNET station and a bird's location.



632

633 **Figure 1.** Comparison of geomagnetic data sources for attaching total intensity (nT) values to a
 634 movement track of a migratory White Stork individual travelling between its wintering ground in
 635 sub-Saharan Africa and its breeding ground in northeastern Europe in spring. **A.** The migratory
 636 track (yellow) of a White Stork individual carrying a GPS tag. INTERMAGNET stations are
 637 identified with black dots. **B.** Difference between using geomagnetic data from nearest
 638 INTERMAGNET station compared to the MagGeo tool output is more evident in locations
 639 where there is a lower density of stations (areas outside of Europe). The fused model estimates
 640 and satellite residual data framework alongside the nearest neighbour interpolation method of
 641 MagGeo ensures a high likelihood of representing the gradient of values experienced by the bird
 642 as it moves across long distances.

643

644 **4. Discussion**

645 The overall absolute difference error was less than 1% of the possible range of values for
646 all geomagnetic components. For example, globally, the total intensity ranges from 23,000 nT to
647 60,000 nT and the mean and median error were 128nT and 80nT respectively when using the fused
648 model estimate and satellite residual data (MG) framework and NNT interpolation method. While
649 animal sensitivity to geomagnetic values is unclear, these error ranges are well within the possible
650 fluctuations of geomagnetic values during geomagnetic storms (>1000 nT for total intensity) and
651 would also be useful at the assumed sensitivity levels for long-distance migratory animals (Beason
652 & Semm, 1987; Chernetsov et al., 2017; Schwarze et al., 2016; Semm & Beason, 1990). Absolute
653 difference stayed below 1% for all geomagnetic components despite changes to the underlying
654 spatiotemporal interpolation method or geomagnetic data structure. NNT did consistently capture
655 the values and patterns observed at ground observatories better than all other interpolation methods
656 thus having the best overall performance. Generally, MG also captured the patterns observed at
657 stations better than just model estimates (CH) especially during high activity levels. More than
658 50% of all stations have moderate to high correlation between MagGeo outputs and observatory
659 values in addition to overall correlation being very high across all geomagnetic components. In
660 general, MagGeo accuracy is lower at higher latitudes and during geomagnetically active periods
661 where there is greater influence of solar activity.

662 In comparison to the Benitez-Paez et al. (2021) who test MagGeo at three
663 INTERMAGNET stations for 6 days in a single year, we test 109 geomagnetic observatory
664 locations across 7 years. Our results suggest that changing MagGeo's underlying spatiotemporal
665 interpolation method from IDW to NNT while continuing to use fused model estimates and
666 satellite residual data will be an improvement to using only model estimates to represent values

667 experienced by animals near the Earth's surface; thus, proving useful for movement ecology
668 studies that test unique research questions regarding the navigational abilities of migratory animals
669 anywhere in the world.

670 *4.1 Interpolation methods*

671 A persistent challenge for the field of remote sensing and movement ecology is how to
672 annotate dynamic environmental covariate data to a moving object (such as a migratory animal)
673 to best model the landscape in a way that accurately represents the animal's experience (Brum-
674 Bastos et al., 2021). This question is addressed by platforms like Movebank (Kays et al., 2022;
675 Kranstauber et al., 2011) and Env-DATA (Dodge et al., 2013) that annotate a movement track with
676 dynamically changing covariates like wind. Env-DATA offers the user with some flexibility for
677 how to interpolate covariate data to the location and timestamp of interest. While the Env-DATA
678 database currently stores information for many different environmental covariates useful for
679 movement ecologists, it does not provide an avenue for attaching geomagnetic data to a movement
680 track.

681 When Benitez-Paez et al., (2021) developed MagGeo to address this technological gap,
682 they implemented an IDW method to interpolate residuals from satellite data to an animal tracking
683 fix. The assumption was that an average geomagnetic value will reduce the influence of any outlier
684 value on the final geomagnetic outputs (Benitez-Paez et al., 2021). However, compared to any
685 nearest neighbour algorithm, averaging the geomagnetic values through IDW may smooth over
686 the very fluctuations MagGeo hopes to capture. We tested this assumption by testing three simpler
687 nearest neighbour interpolation techniques within the MagGeo framework. We found that the NNT
688 had the best performance since it had the highest percent occurrence of lowest error when
689 compared to values and patterns observed at terrestrial geomagnetic stations part of the

690 INTERMAGNET network. Attaching residuals from satellite data closest in time to the point of
691 interest like an animal tracking fix will increase the likelihood of capturing the temporal dynamics
692 of the geomagnetic field. The NNT method is also computationally simpler than IDW algorithm.
693 Based on these results and rationale, there is a strong argument for changing MagGeo's underlying
694 spatiotemporal interpolation from IDW to NNT.

695 It is important to highlight that while we tested different interpolation methods, we
696 maintained the time and space parameters of the existing MagGeo kernel. Benitez-Paez et al.
697 (2021) selected these parameters based on the structure of the polar-orbiting Swarm satellites that
698 pass near a location every four hours with higher clustering of data points at polar latitudes
699 compared to equatorial latitudes due to the radius of the Earth (Friis-Christensen et al., 2006; Olsen
700 et al., 2010). It is important to maintain some function that ensures that temporally
701 contemporaneous geomagnetic data are within a certain spatial distance of an animal tracking fix.
702 Future research may look at what this optimal distance may be to maximize performance, but it is
703 likely that it will vary by latitude like the current implementation.

704 *4.2 Data structure*

705 We found that a combination of CHAOS-7 model estimates and Swarm satellite residuals
706 (MG) can capture relative geomagnetic patterns (lower alpha) observed on the Earth's surface
707 better than using only CHAOS-7 (CH) model estimates. Interestingly, while the satellite data
708 contribution increased as geomagnetic activity increased, this additional residual contribution did
709 not significantly improve the accuracy when comparing with INTERMAGNET stations values
710 (average performance around 50%). This result is surprising since Benitez-Paez et al. (2021)
711 proposed the MG data structure with the expectation that the addition of satellite residuals would

712 capture the temporal variability of the geomagnetic field during periods of higher geomagnetic
713 activity.

714 Model estimates, which are built primarily from quiet-time data, do not represent
715 geomagnetic values during periods of higher activity such as daytime fluctuations outside of quiet-
716 time or discrete events like geomagnetic storms (Finlay et al., 2020; Thébault et al., 2010).
717 Interpolation of residuals from satellites is then likely to add some of the temporal variability not
718 captured by models (Benitez-Paez et al., 2021). Our results indicate however that overall mean
719 MagGeo error increases during periods of higher geomagnetic activity though a positive skew
720 index suggests that mean error may be influenced by a few exceptionally high error data points.
721 Indeed, a double exponential (Laplacian) distribution with a sharp peak around zero and long tail
722 is expected when comparing geomagnetic model data with ground-based measurements (Walker
723 & Jackson, 2000). For example, when comparing IGRF-13 values against ground-based
724 measurements, Beggan (2022) reported absolute difference error values closely mirroring our
725 results alongside the expected Laplacian distribution.

726 Additionally, high geomagnetic activity may impact satellite data collection which will
727 consequently impact the satellite residuals used for the MG structure (Babayev et al., 2006; Lanza
728 & Meloni, 2006). For example, we had to remove data from days during known solar storms as
729 the error was particularly high. Additionally, though we used the NNT interpolation method, the
730 residuals from satellite data may reflect values almost 4 hours before or after the geomagnetic
731 storm due to the MagGeo's space-time cylinder parameters (Benitez-Paez et al., 2021). As a result,
732 while the INTERMAGNET station may have recorded geomagnetic values during the
733 geomagnetic storm, MagGeo outputs may not have captured the localised storm especially at high
734 northern latitudes at such a fine temporal resolution given a possible lagged satellite residual.

735 When combining satellite data and model estimates, we corrected for the difference
736 between satellites collecting geomagnetic data at orbit altitudes and animals experiencing the
737 geomagnetic field at ground altitudes by using model estimates for ground altitudes. The satellite
738 residuals we used however still represent values at orbit altitude and neither our current MagGeo
739 data framework nor interpolation method cannot address this limitation. It is possible to access
740 other geomagnetic model data which, unlike the CHAOS-7 model, provides estimates for all
741 sources of the geomagnetic field. These models do not require correction for the altitude difference
742 (Chulliat et al., 2020; Sabaka et al., 2018, 2020). These values however will not accurately
743 represent the instantaneous and local variability of the geomagnetic field as may be experienced
744 by migratory animals. Incorporating satellite data at satellite altitude will capture some of this
745 variability, and such local variations, especially during geomagnetic storms, are likely to impact
746 orientation by animals using the geomagnetic field (Alerstam, 1987; Benitez-Paez et al., 2021;
747 Schiffner et al., 2011; Wynn et al., 2022; Zein et al., 2022). Further work will be necessary to
748 compare error and accuracy of different models with the MagGeo framework.

749 During most time periods, our results suggest that MG data can consistently capture
750 geomagnetic patterns well. Indeed, while geomagnetic storm events are uncommon, our results
751 suggest that MG still captures geomagnetic patterns slightly better than CH. Additionally, our
752 accuracy analysis uses data from INTERMAGNET stations collected by well calibrated
753 instruments that have high accuracy and precision especially compared to animals who may sense
754 the geomagnetic field for navigation and orientation purposes (Kerridge, 2001; Mouritsen, 2014;
755 Zein et al., 2021). Furthermore, one of the leading hypotheses for animals using the geomagnetic
756 field is the “gradient hypothesis” where relative patterns in geomagnetic intensity or inclination
757 might be more useful for migratory animals than absolute values (Boström et al., 2012; Kishkinev

758 et al., 2015; Wiltschko & Wiltschko, 2021). Additionally, when using the MG data structure and
759 NNT interpolation method, station-wide correlation is moderate to high for most stations. The
760 overall global correlation between MagGeo outputs and observatory values however, is very high
761 suggesting that MagGeo could be used to study long-distance animal migration which happens on
762 a global scale. Based on these considerations, the algorithm from MagGeo that we implement
763 provides a highly useful and robust framework for attaching geomagnetic data to animal tracking
764 data (Zein et al., 2022).

765 *4.3 Outliers*

766 We found outliers to be particularly informative as they highlighted spatial nuances of the
767 geomagnetic field. Most outliers in our dataset represent locations with unique geomagnetic
768 signatures due to local geophysical properties not captured by model estimates at such an acute
769 scale (e.g. Beggan, 2022). Many of the stations highlighted in Figure S1 are volcanic islands with
770 basaltic composition including high concentrations of ferromagnetic minerals (Johnston, 1989;
771 Thébault et al., 2010). Others like the Bangui magnetic anomaly relate to deep geological
772 structures (Girdler et al., 1992). Such lithospheric anomalies have a large local influence on
773 geomagnetic values which may subsequently impact an animal moving through this geomagnetic
774 landscape. For example, birds passing over the geomagnetic anomaly in Sweden have been
775 previously noted to change their behavior suggesting that animal movement may in fact be
776 influenced by local anomalies (Alerstam, 1987). Similar analyses can be conducted by using open-
777 source resources like the World Digital Magnetic Anomaly Map (WDMAM) which allows users
778 to easily extract anomaly information from a raster layer (Lesur et al., 2016).

779 As expected, stations located at high northern latitudes consistently exhibit outliers. In
780 general, all geomagnetic components have a larger range of error at the polar latitudes since

781 charged particles ejected from the sun more readily enter the Earth's atmosphere in the auroral
782 zones at the poles (Campbell, 2003). Model estimates, created from quiet-time data, do not capture
783 these changes for any geomagnetic component. However, for long-distance migrants especially
784 near polar latitudes, geomagnetic strategies may not be useful for navigation or orientation as
785 values can be unreliable in both magnitude and sign. The lack of predictability would thus provide
786 little useful information, especially for migratory animals who have high site fidelity (Lohmann et
787 al., 2008; Wynn et al., 2020). Nevertheless, MagGeo's high global accuracy can still be a valuable
788 tool to reliably attach geomagnetic data to animal tracks for studies wishing to test hypotheses
789 specific to this geographic region.

790 *4.4 Limitations*

791 Geomagnetic sensitivity and perception ranges are unknown for most species and to our
792 knowledge, there are no instruments that accurately record geomagnetic conditions as experienced
793 by migratory animals, though there are species-specific estimates (Åkesson et al., 2005; Beason &
794 Semm, 1987; Chernetsov et al., 2017; Schwarze et al., 2016; Semm & Beason, 1990). The most
795 suitable candidate for attaching geomagnetic data to animal movement data would be
796 INTERMAGNET stations which collect high-temporal resolution geomagnetic values at ground
797 altitudes. These stations however do not have a high global density and thus cannot be used to
798 accurately capture the range of geomagnetic values experienced by an animal during long-distance
799 migration. Using GPS tracking data of a White Stork individual, we demonstrate how attaching
800 geomagnetic data from the nearest station to a migratory bird's location might be ideal for locations
801 in Europe. Outside of Europe however, there would likely be a large mismatch between the
802 geomagnetic values experienced by a bird and a station collecting geomagnetic data more than
803 2,000 km away. Instead, using a combination of model estimates and interpolated satellite

804 residuals could serve as a sufficient alternative that captures high spatiotemporal resolution
805 geomagnetic data for all locations on Earth.

806 We do use INTERMAGNET station data as the ideal standard to perform our error and
807 accuracy analysis to test our MagGeo tool. We did not however anticipate the level of uncertainty
808 introduced by the station data themselves though this is primarily explained by local crustal fields
809 (Beggan, 2022). Our analyses suggest that MagGeo outputs and observatory values are offset by
810 a unique amount specific to each INTERMAGNET station and our linear mixed-effect models
811 reveal that the majority of the error structure for absolute difference can be explained by these
812 random, location-specific effects (St-Louis, 2012). In addition to geomagnetic activity as a fixed
813 effect, these models explain most of the variation in the error structure for differences between
814 MagGeo outputs and observatory values. These results highlight the limitation of our structural
815 set-up as this station-specific offset skews the absolute difference by a consistent amount for each
816 data record. The alpha measure partly addresses this issue by subtracting the standardized MagGeo
817 outputs from the standardized observatory values (Ridley & Macmillan, 2014). Our linear mixed-
818 effect models fit with alpha as the dependent variable suggest that random, location-specific effects
819 explained much less of the error structure. It is noteworthy however that Kp and time of day
820 influenced the error structure in predictable ways such that periods of high geomagnetic activity
821 lead to higher error (Campbell, 2003; Lanza & Meloni, 2006).

822 *4.5 Applications and open questions*

823 Most of our analysis is from data collected in the last 7 years (2014-2020) which is largely
824 during the quiet period of the current solar cycle (Kakad et al., 2020; Li et al., 2011). The 11-year
825 and 22-year solar cycle has a significant influence on geomagnetic field activity since years of
826 high solar activity correspond to higher occurrence of geomagnetic storms (Cliver, 1994; Li et al.,

827 2011; Thébault et al., 2010). Given that CHAOS-7 model estimates contribute three of the four
828 geomagnetic sources (core, lithosphere, and magnetosphere) in MagGeo's framework, we can
829 assume that MagGeo will capture long term changes in the geomagnetic field so long as the
830 CHAOS-7 data inputs are updated. Models capture temporal changes related to secular variation,
831 which arises from changes in the geomagnetic field over a few years due to the motion in the
832 Earth's liquid outer core as well as the slow solar cycle variation of the magnetospheric field
833 (Campbell, 2003).

834 Currently, we are using the CHAOS-7 model estimates (Finlay et al., 2020) but MagGeo's
835 algorithm allows for integration of any other geomagnetic data sources within the VirES platform
836 and may be modified as per the user's need. Specifically, the next couple years of high
837 geomagnetic activity might be of interest to researchers studying the impact of geomagnetic
838 activity on animal behavior. High geomagnetic activity events present a natural occurrence of an
839 experimental extreme that could answer fundamental questions about animal behavior outside of
840 laboratory settings through new "laboratories-in-the-wild" experimental approaches (Nathan et al.,
841 2022). For all above scenarios, the MagGeo tool can facilitate exploration of these research
842 questions.

843 **5. Conclusion**

844 With its relatively low error and flexible framework, MagGeo is a promising tool for
845 movement ecologists and biologists who want to test animal navigation hypotheses about
846 geomagnetism using open, high spatiotemporal resolution geomagnetic datasets. In addition to
847 highlighting the strengths of MagGeo, our study also showcases the importance of error and
848 accuracy tests for environmental covariate data that can be attached to animal movement data. As
849 access to remotely sensed environmental data increases, it will be imperative to enlist cross-
850 discipline expertise to maximize a dataset's full potential and understand the respective strengths
851 and weaknesses of different datasets. Further, our research highlights the need for continued
852 development of analytical tools for combining animal tracking with environmental data. As a
853 research community, we can continue to learn how to better integrate multiple data sources to
854 understand how an animal interacts with its environment thereby contributing to better protections
855 of resources and its inevitable ties to the living world.

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