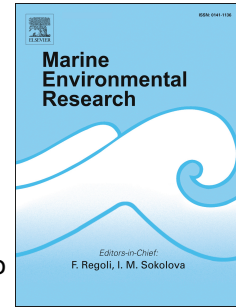


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1 **Physiological response of the coralline alga *Corallina officinalis* L. to both**
2 **predicted long-term increases in temperature and short-term heatwave events**

3
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ABSTRACT

17
18 Climate change is leading to an increase of mean sea surface temperatures and extreme heat events.
19 There is an urgent need to better understand the capabilities of marine macroalgae to adapt to these
20 rapid changes. In this study, the responses of photosynthesis, respiration, and calcification to
21 elevated temperature in a global warming scenario were investigated in the coralline alga *Corallina*
22 *officinalis*. Algae were cultured for 7 weeks under 4 temperature treatments: (1) control under
23 ambient-summer conditions (C, ~20°C), (2) simulating a one-week heatwave of 1°C (HW,
24 $T_{\text{control}+1^{\circ}\text{C}}$), (3
25) elevated temperature (+3, $T_{\text{control}+3^{\circ}\text{C}}$), (4) combination of the two previous treatments (HW+3,
26 $T_{+3+1^{\circ}\text{C}}$). After exposure at T_{+3} (up to a T_{max} of ~23°C), respiration and photosynthesis increased
27 significantly. After 5 weeks, calcification rates were higher at elevated temperatures (T_{+3} and $T_{\text{HW}+3}$)
28 compared to T_{control} , but at the end of the experiment (7 weeks) calcification decreased significantly
29 at those temperatures beyond the thermal optimum (six-fold at T_{+3} , and three-fold at $T_{\text{HW}+3}$,
30 respectively). The same trend was noted for all the physiological processes, suggesting that a
31 prolonged exposure to high temperatures (7 weeks up to T_{+3}) negatively affect the physiology of *C.*
32 *officinalis*, as a possible consequence of thermal stress. A one-week heatwave of +1°C with respect
33 to T_{control} (at T_{HW}) did not affect respiration, photosynthesis, or calcification rates. Conversely, a
34 heatwave of 1°C, when combined with the 3°C increase predicted by the end of the century (at
35 $T_{\text{HW}+3}$), induced a reduction of physiological rates. Continued increases in both the intensity and
36 frequency of heatwaves under anthropogenic climate change may lead to reduced growth and
37 survival of primary producers such as *C. officinalis*.

38
39 **Keywords:** *algae, climate change, ocean warming, temperature, heatwaves, thermal stress,*
40 *calcification, photosynthesis, respiration.*

41

1. Introduction

42

43 Climate change is occurring at a faster rate than in the past, due to increasing concentrations
44 of greenhouse gases in the Earth's atmosphere caused by human combustion of fossil fuels
45 and deforestation (IPCC, 2014). This results in increasing seawater temperatures, rising sea
46 levels, and ocean acidification (IPCC, 2014). The International Panel on Climate Change
47 (IPCC) indicated that global mean surface temperatures have already risen by
48 approximately 0.87°C in the last one and a half centuries (over the period 1850-2015), and
49 will likely increase further (by ca. 3°C by the end of this century, according to the pathways
50 reflecting present nationally stated mitigation goal by 2030; Masson-Delmotte et al., 2018).
51 In addition to long-term warming, extreme events (i.e., storms, droughts, floods and
52 heatwaves) are also becoming more frequent and more intense (Coumou and Rahmstorf,
53 2012; Perkins et al., 2012; Oliver et al., 2018, Frölicher et al., 2018; Darmaraki et al., 2019).
54 Specifically, marine heatwaves (MHWs) can strongly influence ecosystem structure and
55 functioning by causing widespread mortality, species range shifts and community changes
56 (Jentsch et al., 2007; Hobday et al., 2016, Smale et al., 2019).

57 Increasing sea surface temperature (SST) is among the main impacts affecting marine
58 ecosystems (Stenseth et al., 2002), which can influence the abundance and distribution of
59 marine organisms, and lead to poleward range shifts or extinctions of populations located at
60 the edge of their thermal tolerance (Perry et al., 2005; Wernberg et al., 2011; Yara et al.,
61 2012; Jueterbock et al., 2013; Sanford et al., 2016; Collin et al., 2018; Kolzenburg et al.,
62 2019). Elevated temperatures can lead to severe ecological impacts, including widespread
63 mortality of benthic communities (Garrabou et al., 2009), loss of seagrass habitats (Marbà
64 and Duarte, 2010), and impacts on fisheries, due to changes in primary productivity and
65 shifts in distribution or mass mortality events of species of commercial interest (Sumaila et

66 al., 2011; Mills et al., 2013; Caputi et al., 2015). In particular, water temperature is a major
67 factor controlling the survival, growth and reproduction of macroalgae, and thus plays an
68 important role in governing both the small-scale vertical and the large-scale geographical
69 distribution of macroalgal species, in addition to their abundance (Breeman, 1988; Lüning,
70 1990; Nannini et al., 2015). For this reason, it is worth understanding the biological
71 responses of climate-sensitive organisms to short-term extreme events, in concurrence with
72 long-term changes (Jentsch et al., 2007).

73 Coralline red algae (Rhodophyta) are fundamental calcifying primary producers and
74 important habitat-forming species present in most coastal ecosystems, such as coralligenous
75 bioconstructions (Johansen, 1981; Ferrigno et al., 2017; Ingrosso et al., 2018). The species
76 *Corallina officinalis* (Linnaeus 1758) is an erect calcifying alga with a wide distribution that
77 dominates North Atlantic rocky shores and rock pools (Williamson et al., 2015). Due to its
78 complex morphological structure, it represents an important substratum for the settlement of
79 other macroalgae and microalgae, and supports a high biodiversity of marine invertebrates
80 (Akioka et al., 1999; Kelaher, 2003). Despite the importance of coralline algae, their
81 sensitivity to increasing temperatures is still unclear, as different studies have yielded
82 conflicting results (Martin et al., 2013; Comeau et al., 2014; Vásquez-Elizondo and
83 Enríquez, 2016). There is further uncertainty around the response of *C. officinalis* in rock
84 pool habitats, as the species must adapt to multiple stressors, including highly variable water
85 temperatures across seasonal, diurnal and tidal cycles (Williamson et al., 2017).

86 In this study, we describe the physiological responses of *C. officinalis* to temperature
87 variation in an ocean warming scenario (RCP 8.5; IPCC, 2014), taking into account natural
88 thermal fluctuations experienced within rock pools across periods of low and high tides (i.e.,
89 $\Delta T = \sim 3.5^{\circ}\text{C}$, recorded in the field; see Fig. 2). Specifically, we analysed photosynthesis,

90 respiration, and dark/light calcification rates of a South-East UK population exposed to both
91 a temperature increase of +3°C (i.e., simulating the warming expected by the end of this
92 century; Solomon et al., 2007), and a marine heatwave (similar to those registered over the
93 last century, and attributed to anthropogenic climate change; Oliver et al., 2018).

94

95 **2. Materials and methods**

96 *2.1. Biological material*

97 Specimens of the articulated coralline alga *Corallina officinalis* were collected during low
98 tide in intertidal rock pools at ± 0.3 m depth of St. Margarets Bay (Kent, UK;
99 51°08'52.9"N, 1°23'06.9"E) in September 2017. Seawater temperature measured at the time
100 of sampling with a HQ30D flexi multi-meter (Hach Environmental, Loveland, CO, USA)
101 was $15.7 \pm 0.2^\circ\text{C}$. Algae were immediately transported (~3 hours) in temperature-insulating
102 containers to the Institute of Marine Sciences, University of Portsmouth, UK, where the
103 experiment was carried out. Healthy thalli in the size range of 3-10 cm² were selected for
104 the experiment, and were carefully cleaned of epiphytic organisms, avoiding any damage.
105 Algae were fixed on small stones, in order to simulate natural conditions and keep them
106 upright, and guarantee the same light conditions to each branch as much as possible (~3 g
107 fresh weight for each stone), see Fig. 1.

108

109 *2.2. Experimental design*

110 Temperature and irradiance during the experiment were set according to ambient summer
111 conditions recorded daily in the field in July-August 2017 by a HOBO pendant
112 temperature/light data logger (Onset Computer Corp., Bourne, MA, USA) placed in a rock
113 pool of the collection site. Algae were acclimated in 8 15-L aquaria in a closed seawater

114 system for 2 weeks before the gradual increase of water temperature. During
115 acclimatization, algae were maintained at a 14:10 light:dark photoperiod, with UV light
116 oscillating in the range $20\text{-}30\ \mu\text{mol m}^{-2}\text{s}^{-1}$ (measured at the position of the submerged algal
117 fronds in the experimental tanks, mimicking sun set and sun rise and with controlled
118 dimming during the day as cloud effect), while the temperature (T) changed during 24h
119 from a T_{\min} of $\sim 16.5\pm 0.1^{\circ}\text{C}$ to a T_{\max} of $\sim 20\pm 0.1^{\circ}\text{C}$, around a mean temperature value of
120 $18.5\pm 1.2^{\circ}\text{C}$, reflecting the T oscillations of daily tides recorded in the field. After
121 acclimatization, specimens were assigned to 16 x 11-L glass aquaria (4 tanks per treatment)
122 under 4 temperature conditions. Each aquarium contained three algae-stones, for a total of
123 12 algae/stones per temperature treatment (Fig. 1). Aquaria were kept in a closed system
124 with seawater sourced directly from the sea off the Institute of Marine Sciences (University
125 of Portsmouth, UK; $50^{\circ}47'40.7''\text{N}$, $1^{\circ}01'50.1''\text{W}$) and processed via a settlement system
126 with glass media filtration (salinity ranging from 34.4 to 35.2). Ten percent of the aquaria
127 water was exchanged every other day, in order to keep nutrient levels and alkalinity
128 constant. The four temperature treatments (Fig. 2) were: (1) control treatment (C), kept at
129 the *in situ* acclimatization temperature, with a temperature T_{C} oscillating according to a
130 thermal range reflective of the daily tides recorded in the field ($16.5^{\circ}\text{C}\leq T_{\text{C}}\leq 20^{\circ}\text{C}$); (2)
131 heatwave treatment (HW), where a heatwave was simulated by inducing a temperature
132 increase of $+1^{\circ}\text{C}$ for a period of 1 week ($T_{\text{HW}}= T_{\text{C}}+1^{\circ}\text{C}$; $17.5^{\circ}\text{C}\leq T_{\text{HW}}\leq 21^{\circ}\text{C}$); (3) elevated
133 temperature treatment (+3), where the temperature was increased by $+3^{\circ}\text{C}$ according to the
134 predicted temperature increase due to climate change by the year 2100 ($T_{+3}=T_{\text{C}}+3^{\circ}\text{C}$;
135 $19.5^{\circ}\text{C}\leq T_{+3}\leq 23^{\circ}\text{C}$; Solomon et al., 2007); (4) treatment obtained by the combination of the
136 two previous treatments (HW+3), with a $+4^{\circ}\text{C}$ temperature increase ($T_{\text{HW}+3}=T_{\text{C}}+3^{\circ}\text{C}+1^{\circ}\text{C}$;
137 $20.5^{\circ}\text{C}\leq T_{\text{HW}+3}\leq 24^{\circ}\text{C}$). Temperature was increased at a rate of 0.5°C per day (over a period

138 of 6 days) to reach the +3°C temperature change, and of 0.5°C per hour (over a period of 1
139 hour) when simulating the MHW. Water temperature in all tanks was monitored daily with
140 a HQ30D flexi multi-meter (Hach Environmental, Loveland, CO, USA), and had a
141 continual logging every 15 min with a HOBO pendant data logger (Onset Computer Corp.,
142 Bourne, MA, USA). Irradiance levels were monitored throughout the experiment with a
143 Quantitherm light-meter (QRT-1, Hansatech Instruments, Norfolk, UK). pH and salinity
144 were measured using the HQ30D flexi multi-meter by pH and salinity probes (Hach
145 Environmental, Loveland, CO, USA). Total Alkalinity was measured by potentiometric
146 titration (TitroLine 7000, Schott SI Analytics, Mainz, German) following the SOP6 protocol
147 (Dickson et al., 2007). Measurements were validated against Dickson standard (batch #154).
148 Other parameters of the carbonate chemistry were calculated using the software CO2Sys,
149 EXCEL Macro version 2.1 (Lewis et al., 1998). Water motion and filtration in the aquaria
150 was ensured by a submersible pump (V²PowerPump 800, TMC, London, UK).



151

152 **Fig. 1.** Experimental set-up with four temperature treatments (C, HW, +3, HW+3). Each
153 treatment was performed in a large tank, acting as a water bath, in which four 11-L glass
154 aquaria were immersed (a total of 16 aquaria). Every aquarium contained three algae-stones
155 (as showed in the detail, top right).

156

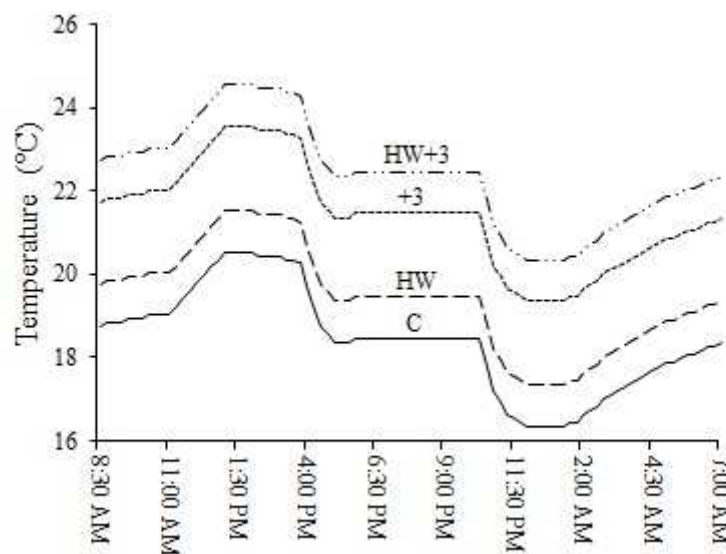
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158

159 *2.3. Marine heatwave calculation*

160 A one-week MHW of 1°C was calculated in accordance to the definition provided by
 161 Hobday et al. (2016), i.e., referring to the temperature values exceeding the 90th percentile
 162 threshold of the SST measured for at least five consecutive days in the same 30-day-period
 163 window over the last 30 years. SST *in situ* data were obtained from the closest NOAA buoy
 164 to the collection site, located off the South UK Coast (about 40 km) along the Greenwich
 165 meridian (Station 50°24'0" N 0°0'0" E; National Data Buoy Center, National Oceanic and
 166 Atmospheric administration; www.ndbc.noaa.gov/station_page.php?station=62305).

167



168

169 **Fig. 2.** Daily planned temperature changes in the experimental tanks for the 4 treatments (C,
 170 HW, +3, HW+3). The temperature fluctuation in the control (C) was performed simulating
 171 the environmental thermal excursion due to the daily tides, as recorded in the field. The
 172 MHW was performed in the treatments HW and HW+3, and lasted for one week (for more
 173 details see Fig. 3).

174

175

176 *2.4. Physiological measurements*

177 Algal thalli (1 g/fresh weight from each tank at each time point) were incubated in 50-ml
 178 closed oxygen chambers filled with bubble-free seawater from the aquaria. One hour
 179 incubations were conducted under saturating light condition ($300 \mu\text{mol m}^{-2}\text{s}^{-1}$; Ralph and
 180 Gademann, 2005) and in the dark. The irradiance levels were controlled with a Quantitherm
 181 Light Meter (QRT-1, Hansatech Instruments, Norfolk, UK). The chambers were used to
 182 assess net photosynthesis (P_n) and calcification in the light (G_l), while chambers covered
 183 with aluminium foils were used to assess dark respiration (R_d) and calcification in the dark
 184 (G_d). The concentration of dissolved oxygen (O_2 , $\mu\text{mol l}^{-1}$) was measured inside the
 185 chambers before and after incubations using a HQ30D flexi oxygen meter (Hach
 186 Environmental). Water samples were taken at the beginning and at the end of the
 187 incubations for measurements of pHT (pH on the total scale) and total alkalinity (A_T).
 188 P_n and R_d , expressed in terms of O_2 production and consumption (in $\mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$),
 189 were calculated after Williamson et al. (2017):

$$P_n(\text{or } R_d) = \frac{\Delta O_2 v}{fw \Delta t}$$

190 where ΔO_2 is the difference in O_2 concentration before and after incubation ($\mu\text{mol l}^{-1} \text{ h}^{-1}$), v
 191 is the volume of the incubation tubes (l), fw is the fresh weight of the algae incubated (g)
 192 and Δt is the incubation time (h).

193 Gross photosynthesis (P_g) was calculated as:

$$P_g = |P_n| + |R_d|$$

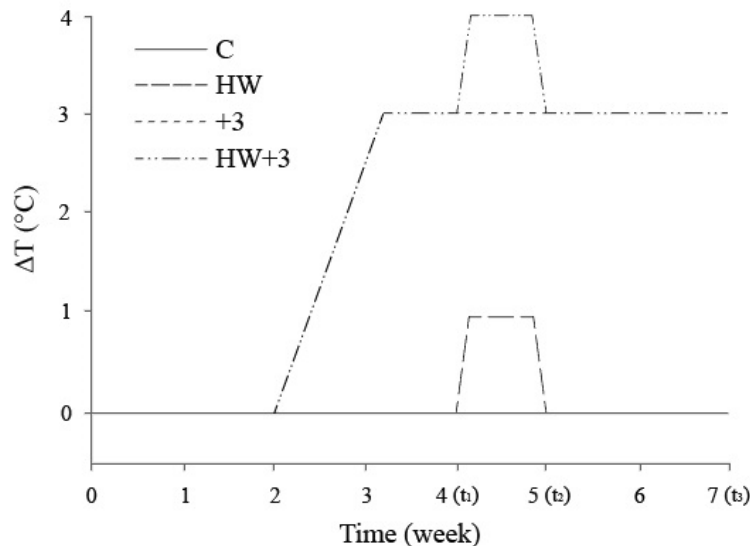
194 G_l and G_d ($\mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$) were calculated using the alkalinity anomaly technique
 195 (Smith and Key, 1975) as:

$$G_l(\text{or } G_d) = \frac{\Delta A_T v}{2(f_w \Delta t)}$$

196 where ΔA_T is the difference between initial and final A_T values ($\mu\text{eq l}^{-1}$).

197 Physiological measurements were taken in all treatments at 3 different times during the
198 experiment (Fig. 3):

- 199 • t_1 = before the MHW start (4-weeks); i.e., after 2 weeks of acclimatization at the initial
200 temperature T_{control} , 1 week of gradual heating up to T_{+3} , and 1 week of acclimatization at
201 T_{+3} (the heating up to T_{+3} was induced only in treatments +3 and HW+3);
- 202 • t_2 = right after the MHW end (5-weeks); i.e., 1 week of HW at T_{HW} after t_1 (the HW was
203 induced only in treatments HW and HW+3);
- 204 • t_3 = after a recovery period from the MHW end (7-weeks); i.e., 2 weeks of recovery from
205 the heatwave-end after t_2 .



206
207 **Fig. 3.** Planned temperature changes during the experiment in the 4 treatments (C, HW, +3,
208 HW+3). Physiological measurements were taken in all treatments at 3 different times: t_1 (4-
209 weeks), before the MHW start; t_2 (5-weeks), right after the MHW end; t_3 (7-weeks), after a
210 recovery period from the MHW end.

211

212 2.5. Data analysis

213 We used linear-mixed effects models (LMMs) to examine whether temperature treatments
214 influenced photosynthesis, respiration, and calcification rates. Models were developed in the
215 *nlme* package in R v3.6.0 (Bates et al., 2015; Pinheiro et al., 2019) with both time and
216 temperature (as well as their interaction) treated as fixed effects, and tank ID included as a
217 random effect to account for autocorrelated errors among algae grown in the same tanks
218 (Speights et al., 2017). Model residuals were visually inspected using QQ plots and residual
219 plots and formally checked for normality and homoskedasticity via Shapiro-Wilks and
220 Levene's tests, respectively. Heterogeneity in residual variance was only identified in the
221 calcification models, and was addressed with an appropriate structure (*varIdent*), allowing
222 residuals to differ in spread between temperature treatments across time without the need to
223 transform the data (Pinheiro and Bates, 2000; Harrison et al., 2018). Results are expressed
224 as mean \pm standard error of the mean (SE). n is the sample size and p_s are Tukey-adjusted p -
225 values, evaluated against a significance threshold of $\alpha = 0.05$.

226 3. Results

227 3.1. Respiration

228 LMMs explained over 76% of the variation in dark respiration (R_d) [conditional pseudo- R^2
229 = 0.763; Nakagawa and Schielzeth, 2013]. At time t_1 , before mimicking the marine
230 heatwave, no significant differences were found between R_d rates in the control (C) and the
231 treatment HW, and between the treatments +3 and HW+3 (Fig. 4; Table 1). This trend was
232 expected because, at time t_1 , the treatments C and HW were kept at the same temperature T_c
233 ($16.5^\circ\text{C} \leq T_c \leq 20^\circ\text{C}$); while, +3 and HW+3 were both at T_{+3} ($T_{+3} = T_c + 3^\circ\text{C}$; $19.5^\circ\text{C} \leq T_{+3} \leq 23^\circ\text{C}$;
234 see fig. 3). However at t_1 , C and HW were both different from +3 and HW+3, with lower R_d
235 values measured in +3 and HW+3 (0.88 ± 0.08 and $0.81 \pm 0.09 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$,

236 respectively), and higher values in C and HW (1.27 ± 0.15 and $1.38\pm 0.12 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$,
 237 respectively) [$p_{C,+3} = 0.049$, $p_{C,HW+3} = 0.021$, $p_{HW,+3} = 0.012$, $p_{HW,HW+3} = 0.005$]. At time t_2 ,
 238 immediately after the MHW, the lowest R_d rates were measured in the HW treatment
 239 ($0.53\pm 0.06 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$), while the highest rates were found in the treatment +3
 240 ($1.22\pm 0.04 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) [$p_{HW,+3} = 0.001$]; intermediate values were measured in C
 241 ($0.90\pm 0.07 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$). At time t_3 , in HW+3, the lowest R_d values were registered
 242 ($0.45\pm 0.10 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) [$p_{C,HW+3} = 0.026$].
 243 R_d in the control group did not change significantly over time (Fig. 4). In the treatment +3,
 244 R_d rates increased from t_1 to t_2 (0.88 ± 0.10 and $1.22\pm 0.04 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$, respectively)
 245 [$p_{t_1,t_2} = 0.023$], and decreased from t_2 to t_3 (1.22 ± 0.04 and $0.88\pm 0.07 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$,
 246 respectively) [$p_{t_2,t_3} = 0.027$]. In HW, R_d rates decreased from t_1 to t_2 , going from 1.38 ± 0.12
 247 to $0.53\pm 0.04 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$ [$p_{t_1,t_2} < 0.001$]. In HW+3, R_d rates decreased from t_1 to t_3
 248 (0.81 ± 0.09 to $0.45\pm 0.10 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) [$p_{t_1,t_2} = 0.015$]. While, no changes between t_2
 249 (right after the MHW end) and t_3 (after the recovery period from the MHW) were observed
 250 in both the heatwave conditions HW and HW+3.

251

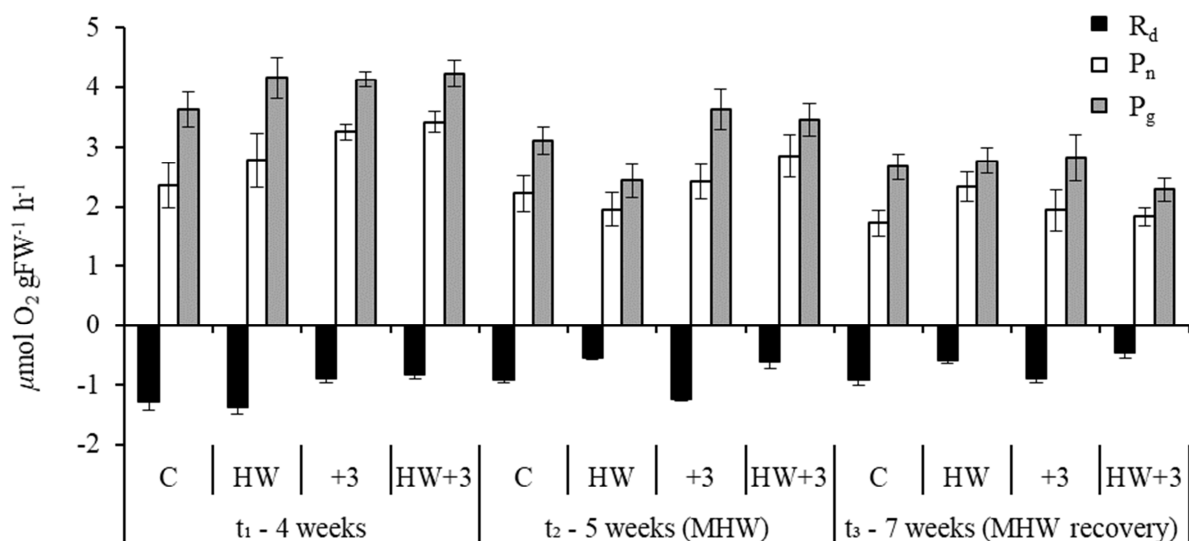
252 3.2. Photosynthesis

253 LMMs explained over 58% of the variation in net photosynthesis (P_n) and 67% of variation
 254 in gross photosynthesis (P_g) [conditional pseudo- $R^2 = 0.581$ and 0.672 , P_n and P_g
 255 respectively]. No differences in P_n rates were found among temperature treatments at each
 256 experimental time (t_1 , t_2 , t_3 ; Fig. 4; Table 1). Likewise, no differences in P_g rates were
 257 apparent at time t_1 between the control (C) and the treatment at $+3^\circ\text{C}$ (3.64 ± 0.29 and
 258 $4.13\pm 0.12 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$, respectively), or between HW and HW+3 (4.16 ± 0.35 and

259 $4.23 \pm 0.23 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$, respectively). At time t_2 , the lowest P_g values were measured
 260 in the treatment HW ($2.44 \pm 0.28 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) and the highest in +3 (3.64 ± 0.33 and
 261 $\mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$), with these being significantly different from each other [$p_{\text{HW},+3} =$
 262 0.036]. No significant differences were found among all temperature treatments at time t_3 ,
 263 after the recovery period from the MHW end.

264 In C, P_n rates did not vary among experimental time points (t_1 , t_2 and t_3), while P_g decreased
 265 from t_1 to t_3 [$p_{t_1,t_3} = 0.028$]. In the treatment +3, both P_n (from 3.26 ± 0.13 to $1.94 \pm 0.35 \mu\text{mol}$
 266 $\text{O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) and P_g (from 4.13 ± 0.12 to $2.83 \pm 0.38 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$) decreased from t_1 to
 267 t_3 [$p_{t_1,t_3} = 0.004$ and 0.003 , P_n and P_g respectively]. In the treatment HW, P_n did not change
 268 significantly overtime, while P_g decreased from t_1 to t_2 (4.16 ± 0.35 and $2.78 \pm 0.22 \mu\text{mol O}_2$
 269 $\text{gFW}^{-1} \text{ h}^{-1}$, respectively) [$p_{t_1,t_2} < 0.001$], as well as from t_1 to t_3 [$p_{t_1,t_3} = 0.001$]. In HW+3, the
 270 lowest P_n and P_g rates were found after the recovery period from the MHW end (t_3 ;
 271 1.84 ± 0.15 and $2.29 \pm 0.20 \mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$, respectively), with t_3 differing from both t_1 and
 272 t_2 [$p_{t_1,t_3} < 0.001$ and 0.001 , $p_{t_2,t_3} = 0.025$ and 0.007 , P_n and P_g respectively].

273



274

275 **Fig. 4.** Net and gross photosynthesis (P_n and P_g) rates at the experimental irradiance, and
276 respiration (R_d) rates in the dark in the 4 temperature treatments (C, HW, +3, HW+3) for the
277 3 incubation times (t_1 , before the MHW start; t_2 , right after the MHW end; t_3 , after a
278 recovery period from the MHW end). Data are means \pm SE, and are expressed in terms of
279 O_2 release (negative values for respiration correspond to O_2 consumption); $n = 4$ for each
280 treatment.

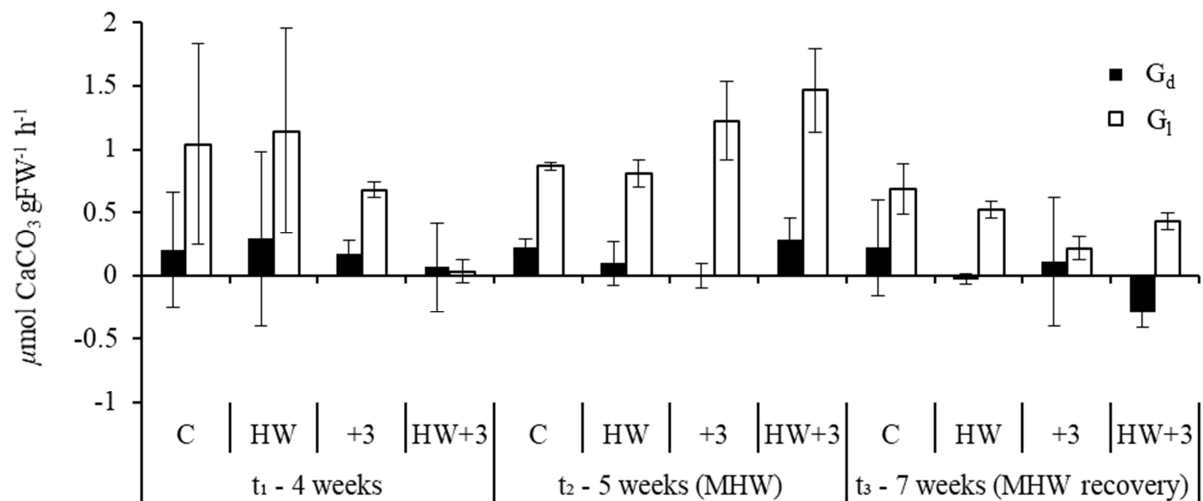
281

282 3.3. Calcification

283 There was only weak evidence for an effect of temperature on calcification rates measured
284 in the dark (G_d ; Table 1), with no post-hoc differences found among temperature treatments,
285 highlighting uncertainty about how temperature affects G_d rates. This is in alignment with
286 the low percentage of variance explained by the LMMs [G_d conditional pseudo- $R^2 = 0.034$].
287 Variability in G_1 rates was inherently high, especially at the start of the experiment (i.e.,
288 time t_1), leading to LMMs with poor explanatory power [conditional pseudo- $R^2 = 0.0601$].
289 G_1 rates were affected by temperature at time t_1 , with +3 and HW+3 being significantly
290 different from each other [$p_{+3,HW+3} < 0.001$], despite being at the same temperature T_{+3} . At
291 time t_2 , +3 and HW+3 (1.22 ± 0.31 and $1.47 \pm 0.33 \mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$) exhibited (non-
292 significant) higher G_1 rates than C and HW (0.87 ± 0.03 and $0.81 \pm 0.10 \mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$,
293 respectively). After 2-weeks-recovery from the MHW (t_3), an opposite trend was
294 observed, with +3 and HW+3 characterized by the lowest G_1 rates (0.21 ± 0.09 and 0.43 ± 0.07
295 $\mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$), and C and HW by the highest rates (0.69 ± 0.20 and $0.52 \pm 0.07 \mu\text{mol}$
296 $\text{CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$, respectively). Although non-significant, at time t_3 the lowest G_d rates
297 were measured in the treatment HW+3, with negative values corresponding to a net
298 dissolution of the algae ($-0.30 \pm 0.11 \mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$), and the highest ones were

299 found in C ($0.22 \pm 0.38 \mu\text{mol CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$, with one sample having experienced
 300 dissolution).

301 G_d rates did not change significantly over time (Table 1). In C and HW, no differences in G_1
 302 rates were observed among experimental time points (t_1 , t_2 and t_3). G_1 rates decreased from t_2
 303 to t_3 in the $+3^\circ\text{C}$ condition, with rates varying from 1.22 ± 0.31 to $0.21 \pm 0.09 \mu\text{mol CaCO}_3$
 304 $\text{gFW}^{-1} \text{ h}^{-1}$ [$p_{t_2, t_3} = 0.001$]; and in the HW+3 treatment, from 1.47 ± 0.33 to $0.43 \pm 0.07 \mu\text{mol}$
 305 $\text{CaCO}_3 \text{ gFW}^{-1} \text{ h}^{-1}$ [$p_{t_2, t_3} = 0.015$].



306
 307 **Fig. 5.** Calcification rates in the dark (G_d) and at the experimental irradiance (G_1) in the 4
 308 treatments (C, HW, +3, HW+3) for the 3 incubation times (t_1 , before the MHW start; t_2 ,
 309 right after the MHW end; t_3 , after a recovery period from the MHW end). Negative values
 310 for algal calcification correspond to the decalcification activity quantified as increase in
 311 total alkalinity. Data are expressed as means \pm SE; $n = 4$ for each treatment.

312

Table 1

Summary of linear mixed effects models testing the effect of temperature and experimental time on *C. officinalis* metabolism in the dark and at the incubation irradiance. R_d , dark respiration; P_n , net production; P_g , gross production; G_d , net calcification in the dark; G_l , net calcification in the light. Bolded values indicate p -values < 0.05 .

Factor	df	Dark				Light					
		R_d		G_d		P_n		P_g		G_l	
		F-value	p -value	F-value	p -value	F-value	p -value	F-value	p -value	F-value	p -value
Temperature	3	8.374	0.003	4.962	0.018	1.568	0.248	1.124	0.378	32.424	<.001
Time	2	21.963	<.001	3.030	0.068	15.261	<.001	33.143	<.001	7.393	0.004
Interaction	6	8.836	<.001	0.833	0.557	1.873	0.127	2.536	0.048	6.733	<.001

4. Discussion

345

346 Macroalgae occurring in intertidal habitats (such as *C. officinalis*) are exposed to strong
347 daily and seasonal temperature fluctuations. As such, they may have a greater ability to
348 acclimate to higher temperature regimes than macroalgae found in more thermally stable
349 conditions. However, little is known of how extreme changes in water temperature, e.g.
350 heatwaves, affect algae physiology in coastal rock pools.

351 Temperature has a fundamental effect on chemical reaction rates, and a general dependence
352 of respiration and photosynthesis to temperature is well known in macroalgae (Lüning,
353 1990).

354 Our results for *C. officinalis* confirmed this dependence by showing, initially at time t_1 , a
355 decrease in respiration rates at elevated temperatures, followed by an opposite trend after the
356 MHW simulation, at t_2 and t_3 , with an increase in respiration rates with higher temperatures.
357 This increase in respiration rates at elevated temperatures is in line with other studies carried
358 out on different species of coralline algae (Adey, 1973; Digby, 1977; Ichiki et al., 2001;
359 Martin et al., 2006; Steller et al., 2007, Williamson et al., 2017). In particular, when the
360 MHW was simulated, we observed an increase in respiration rates at temperatures raised by
361 $+3^\circ\text{C}$ relative to measured summer values (i.e., up to a T_{max} of about 23°C ; see Fig. 2).
362 However, a further 1°C increase (mimicking a MHW) in the HW+3 treatment (i.e., up to a
363 T_{max} of about 24°C) led to respiration rate reductions. This shows that a decline in
364 respiration occurs beyond a thermal optimum that is close to the SST registered in the
365 summer season (i.e., as simulated in our experiment with T_c daily oscillating in the range
366 $16.5\text{-}20^\circ\text{C}$). The same trend is confirmed after 2 weeks of recovery from the MHW end, at
367 time t_3 . These results are in accordance with those reported by Martin et al. (2013) on the
368 temperate coralline alga *Lithophyllum cabiochae*, demonstrating a positive effect on

369 respiration rates at higher temperatures during the colder months, and either a negative or a
370 nil response during the summer when temperatures are closer to a thermal optimum
371 (Anthony et al., 2008). We also observed that a prolonged exposure to high temperatures
372 (i.e., t_3 , 7 weeks up to a T_{\max} of about 23°C in daily temperature variation) negatively
373 affected respiration rates, with the lowest respiration values registered in HW+3 (0.45 ± 0.10
374 $\mu\text{mol O}_2 \text{ gFW}^{-1} \text{ h}^{-1}$).

375 There was limited evidence for an effect of temperature on photosynthesis of *C. officinalis*,
376 despite a significant increase of P_g rates with temperature, reaching a maximum of +3°C
377 variation with respect to the control, was observed after 5 weeks (at t_2). Although non-
378 significant, the same increase is observed in P_n . This general trend mirrors the results
379 reported for *C. officinalis* (Digby, 1977; Williamson et al., 2017) and other coralline algae
380 (Digby, 1977; Ichiki et al., 2001; Martin et al., 2006; Steller et al., 2007; Martin et al.,
381 2013), which indicate higher photosynthesis variation as a consequence of elevated
382 temperature (c.a. 10°C) and irradiance changes between winter and summer (Martin et al.,
383 2013; Williamson et al., 2017). By simulating summer conditions of irradiance and SST in
384 our experiment, we recorded small positive variations in photosynthesis between the control
385 and the elevated temperature treatments. This might occur at temperatures already close to
386 the thermal optimum (Anthony et al., 2008), and partially agrees with the observation of
387 Martin et al. (2013) in *L. cabiochae* where significant effects of the 3°C warming were
388 detected on P_g in colder seasons but not in the summer. Importantly, as already noted for
389 respiration, a prolonged exposure to high temperatures (i.e., t_3 , 7 weeks up to a T_{\max} of about
390 23°C) negatively affected P_n rates, as a possible effect of thermal stress.

391 In general, dark calcification showed high variability in all treatments at all experimental
392 time points. This reflects the findings of Kolzenburg et al. (2019) on *Corallina officinalis*,

393 and suggests that the already small amount of calcification in the dark is easily influenced
394 by environmental factors such as temperature. However, we observed lower rates of calcium
395 carbonate precipitation in the dark with respect to experimental irradiance. This is in line
396 with the results reported for *Amphiroa anceps* and *A. foliacea* by Borowitzka (1981), for
397 *Corallina frondescens* and *C. vancouveriensis* by McCoy et al. (2016), and for *C. officinalis*
398 by Kolzenburg et al. (2019), due to the strict connection between algal photosynthetic
399 activity, providing the greatest contribute to CO₂ fixation, and calcification. The rates of
400 calcification under experimental irradiance and in the dark did not exhibit significant
401 responses to temperature within each time point. Previous studies on coralline algae showed
402 similar results, with high variations in calcification rates reported for coralline algae under
403 high changes of both irradiance and temperature (Martin et al., 2013; McCoy et al., 2016;
404 Williamson et al., 2017; Kolzenburg et al. 2019). In the present study we only considered
405 lower temperature regimes and constant values of irradiance compared to previous studies.
406 However, a general (but non-significant) calcification increase under summer irradiance
407 conditions was observed in the first 5 weeks of the experiment for a +3°C (and also +4°C
408 after 4 weeks) variation with respect to the control temperature. This agrees with the
409 lowering of calcification rates at cooler temperatures seen in other *Corallina* species (*C.*
410 *frondescens* and *C. vancouveriensis*) by McCoy et al. (2016). Critically, the significant
411 decrease measured in G₁ rates at elevated temperatures (six fold for the +3°C condition, and
412 threefold for the HW+3 condition, respectively) at the end of the recovery time (from t₂ to
413 t₃) may possibly be related to thermal stress. These results reflect findings of Vásquez-
414 Elizondo and Enríquez (2016) on the coralline algae *Amphiroa tribulusform*,
415 *Neogoniolithon sp.* and *Lithothamnion sp.*, indicating losses in algal calcification after
416 exposure to elevated temperature (+ 2°C above the local maximum monthly mean

417 temperature). This decreasing trend is similar in the case of dark calcification, with a
418 significant decrease of G_d in the $+4^\circ\text{C}$ condition at the end of the recovery time, consistent
419 with the effect of a stress induced by a prolonged warming.

420 Furthermore, it has to be considered that under anthropogenic climate change, MHWs will
421 likely increase in intensity and frequency (Hobday et al., 2016), and that coralline algae will
422 possibly be affected by other stresses such as ocean acidification (Hall-Spencer et al., 2008;
423 Kuffner et al., 2008; Martin and Gattuso, 2009; Ragazzola et al., 2012; Donnarumma et al.,
424 2014). Our findings indicated that MHWs could be more harmful when combined to the
425 long-term temperature increase predicted by the end of the century. This overall increase in
426 temperature and the increase in frequency and intensity of the heatwaves could thus have
427 severe effects on the species' distribution, creating a range shift northwards (Araújo et al.,
428 2005). Therefore, southern margin populations of *C. officinalis*, together with the high
429 densities of macrofaunal organisms living within their fronds, may risk disappearing in their
430 original environment as temperatures warm in future oceans (Kolzenburg et al., 2019).
431 However, at higher latitudes, aragonite saturation state and ocean pH will reach critically
432 low levels first (Steinacher et al., 2009), potentially leading to a shift of calcifying species
433 distribution southwards (Orr et al., 2005; Yara et al., 2012; Lenton et al., 2015). This might
434 result in a contraction of the natural distribution of *C. officinalis*, yet more studies
435 considering the combined effects of warming and acidification are needed to predict how
436 this species' distribution and abundance may be affected by anthropogenic climate change.

437

438

5. Conclusion

439 By combining the effects of thermal stress induced by tides, gradual ocean warming, and
440 marine heatwaves, we have shown how important insights can be obtained on the likely

441 physiological responses of coralline algae to climate-change induced temperature variations.
442 Our experiment was designed to simulate summer conditions, in order to understand how
443 this species will respond to temperature variations stemming from climate warming when
444 the algae already experience conditions that are near their thermal optimum. Our results
445 indicated that temperature has a significant effect on *C. officinalis* physiology. After
446 exposure to +3°C from the field temperature (up to a T_{\max} of about 23°C), both respiration
447 and photosynthesis increased. This trend has already been reported for several species of
448 coralline algae (Martin et al., 2013). After 5 weeks, calcification seemed to be enhanced at
449 higher temperatures (up to a T_{\max} of about 23 and 24°C), but at the end of the experiment
450 calcification rates decreased at those temperatures beyond the thermal optimum. The same
451 trend was noted for all the physiological processes, suggesting that a prolonged exposure to
452 high temperatures (i.e., 7 weeks up to a T_{\max} of about 23°C) negatively affects the
453 physiology of *C. officinalis*, as a possible effect of thermal stress. A one-week heatwave of
454 +1°C with respect to the control temperature did not significantly affect respiration,
455 photosynthesis, or calcification rates. This might be explained by the good adaptation of *C.*
456 *officinalis* to both seasonal and tidal temperature variability (Williamson et al., 2017).
457 Conversely, a further increase of 1°C (due to the MHW) to the 3°C increase predicted by
458 the end of the century, often induced physiological rate reductions, underlining that MHWs
459 may have a negative impact on this species in the near future. Given the fundamental
460 ecological role of *C. officinalis* and other coralline algae as habitat-forming species, stronger
461 and more frequent temperature extremes over the next decades could result in a decrease in
462 coralline algal abundance or a shift in the species' distribution, with potentially major
463 consequences for biodiversity in coastal ecosystems.

464

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Highlights

- After a 7-weeks exposure, physiological rates decreased at elevated temperatures.
- A one-week heatwave of 1°C did not affect algal metabolism.
- Heatwave of 1°C added to the 3°C increase induced a reduction of algal metabolisms.